

Successes and Future Challenges of Mantle Modelling

Yanick Ricard

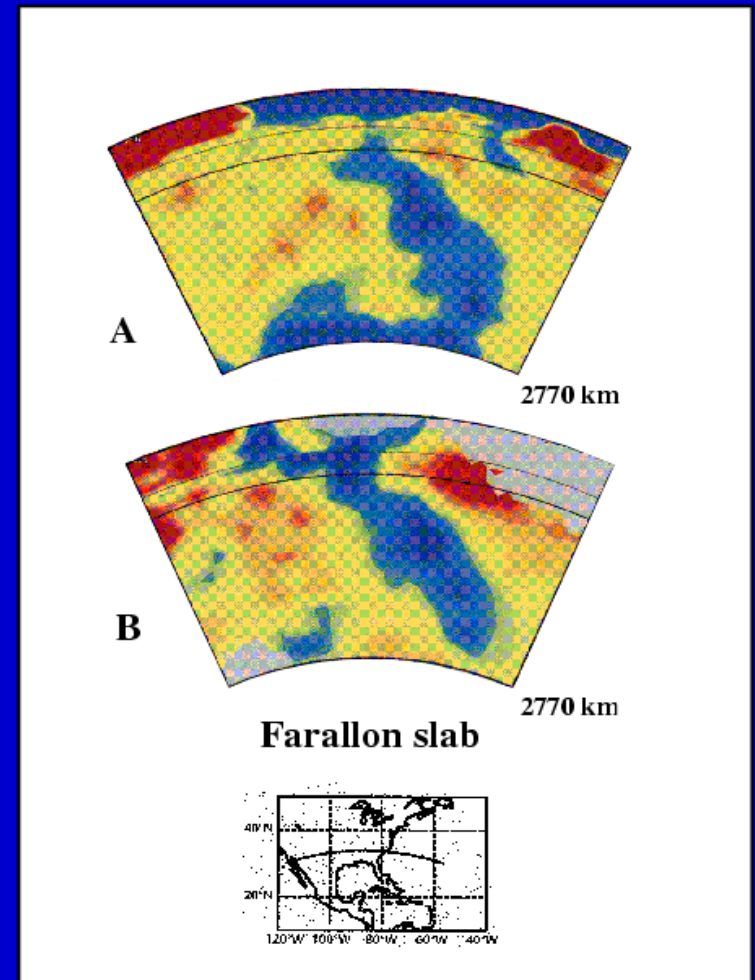
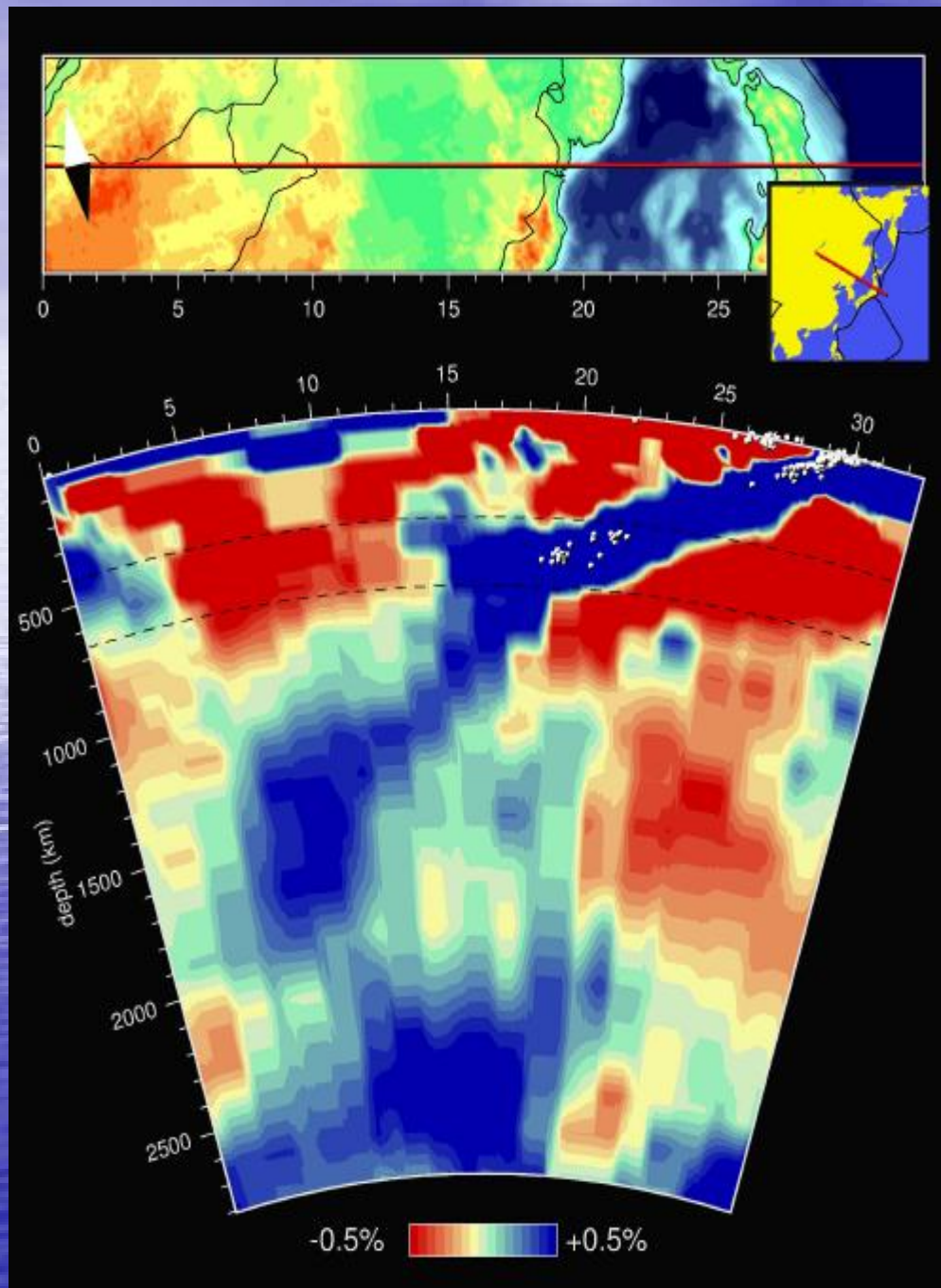
CNRS/ENS-Lyon/Universite-Lyon

1) Density anomalies

2) Rheology

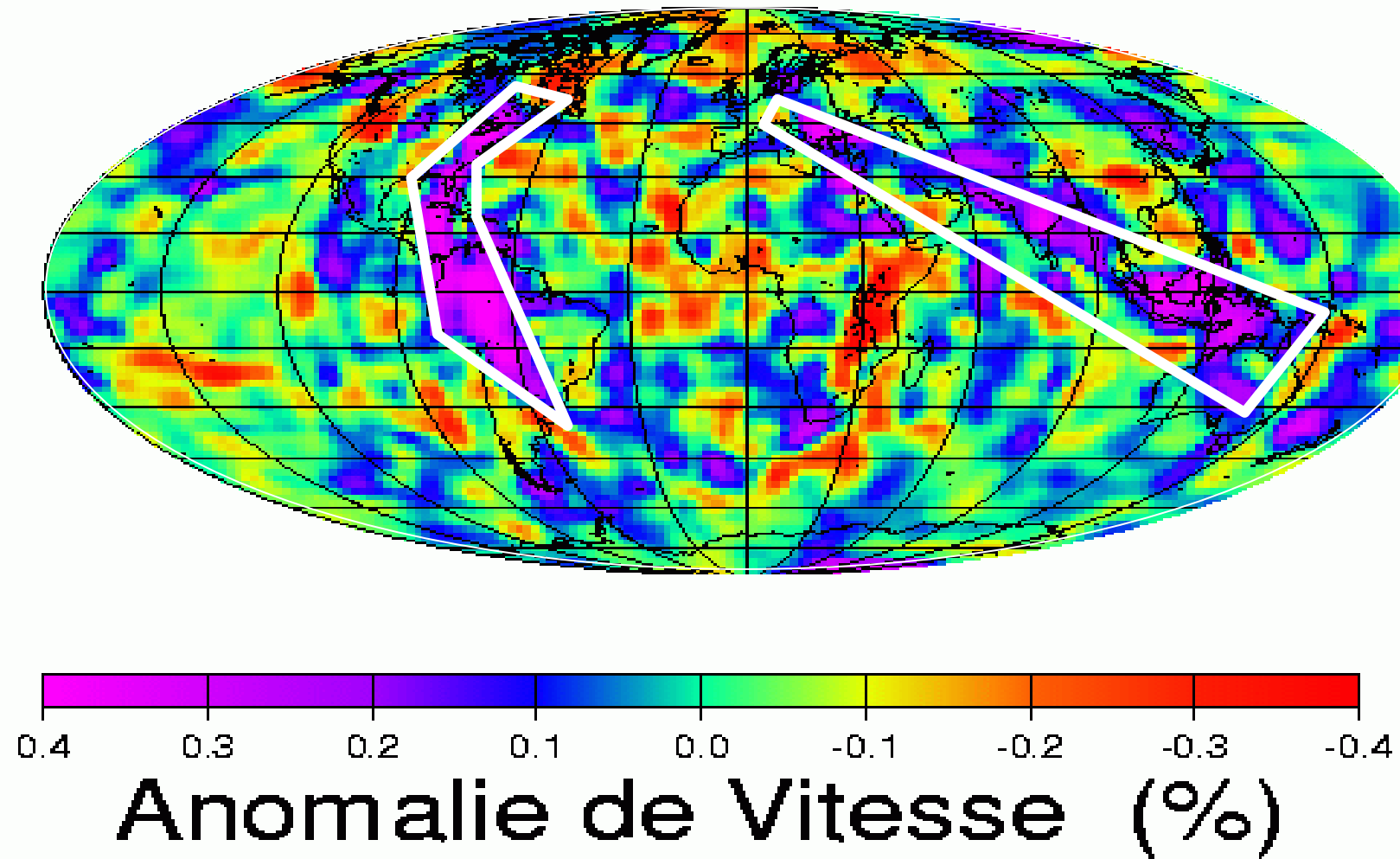
3) Mixing/geochemistry

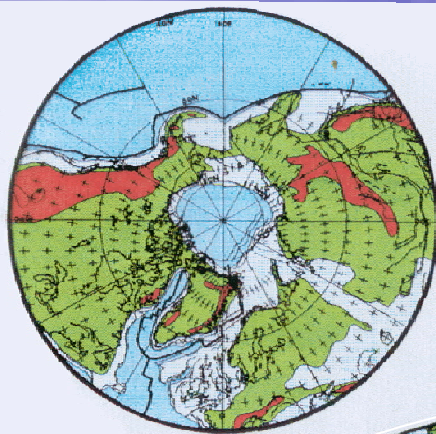
In the last 10 years the links
between plate tectonics and
mantle convection became
much more obvious...



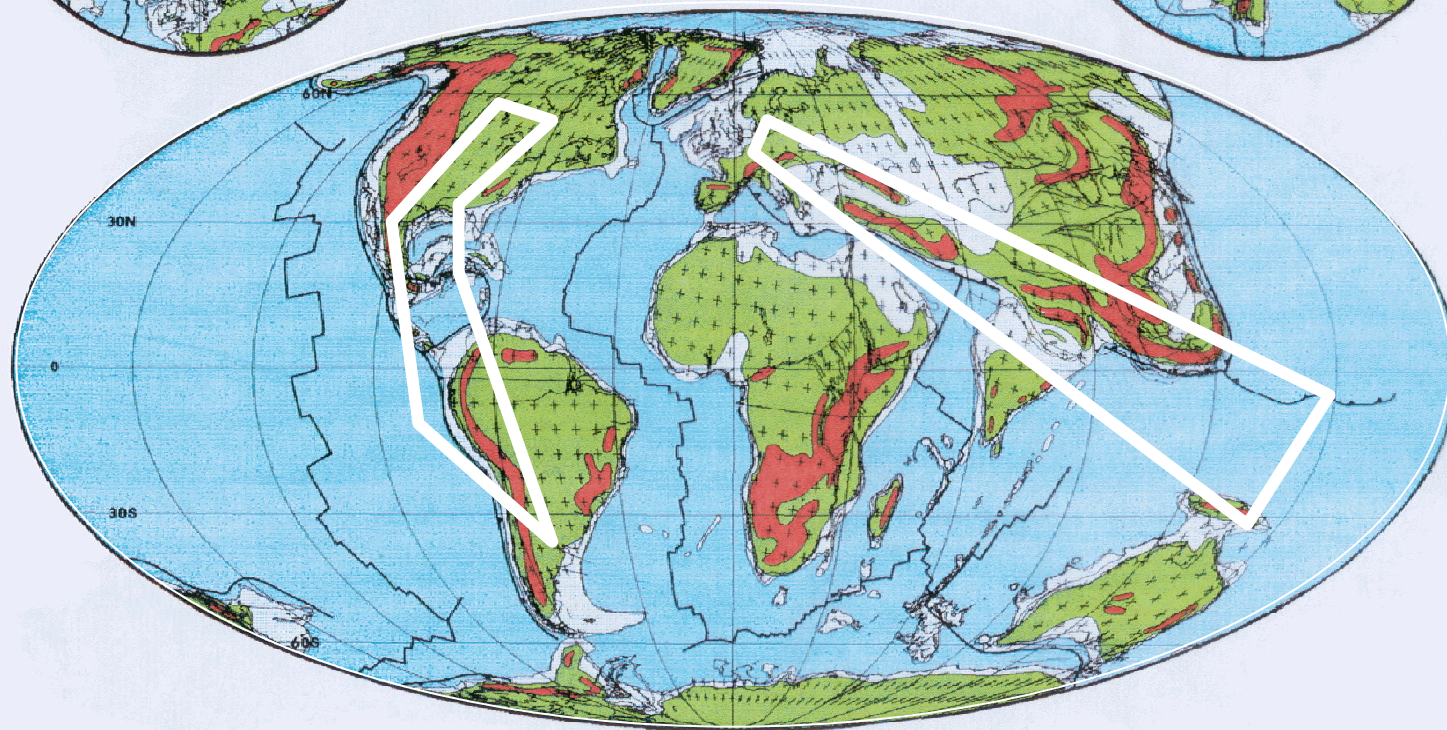
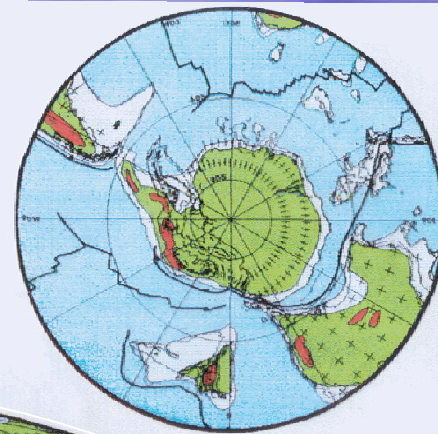
From van der Hilst, Grand, Spakman...

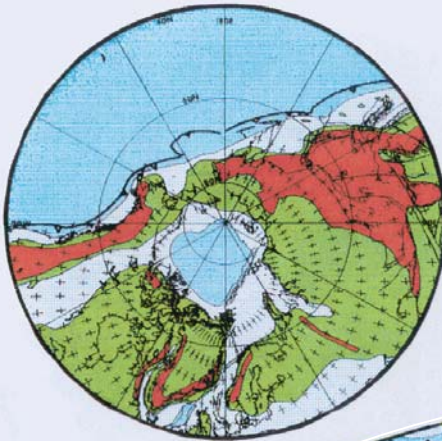
(Grand, 1997) 1100 km



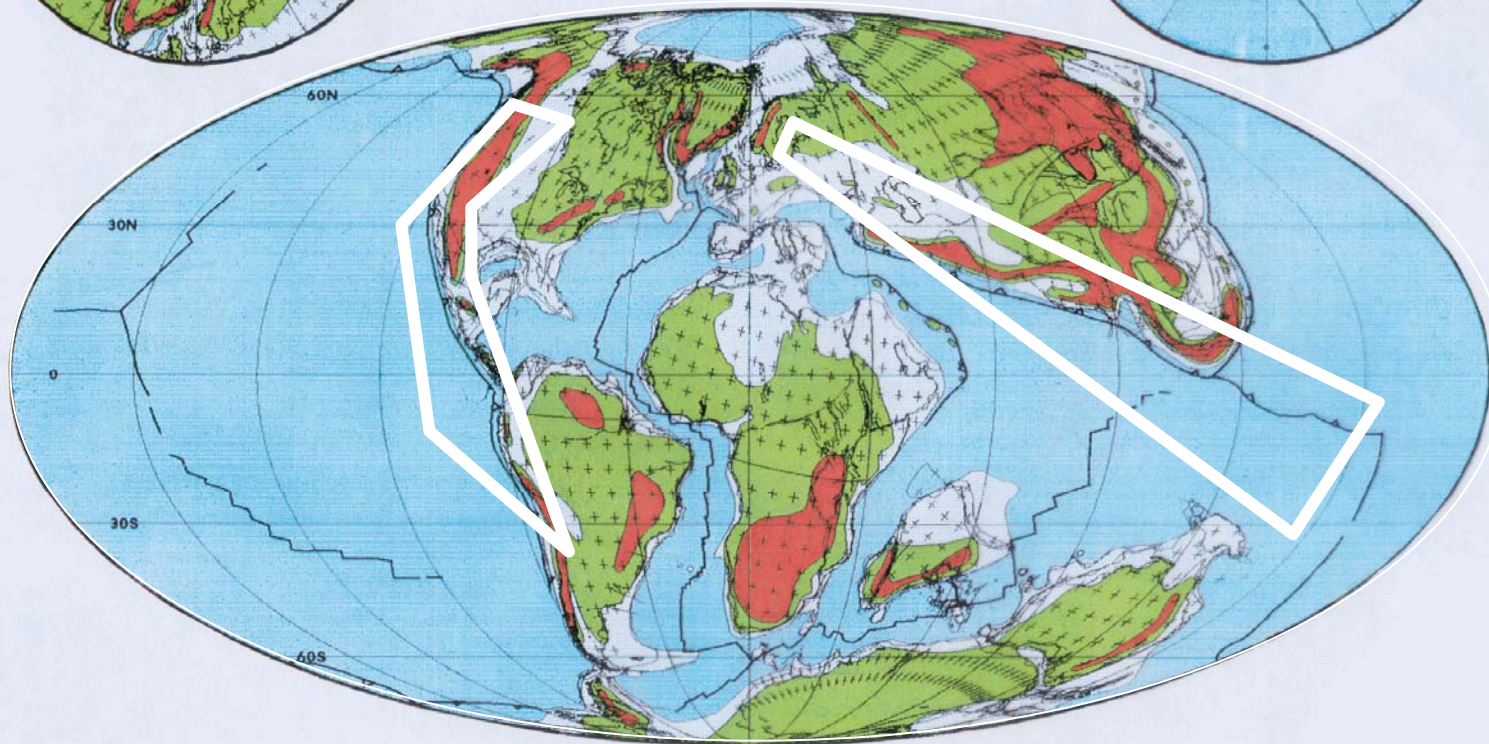
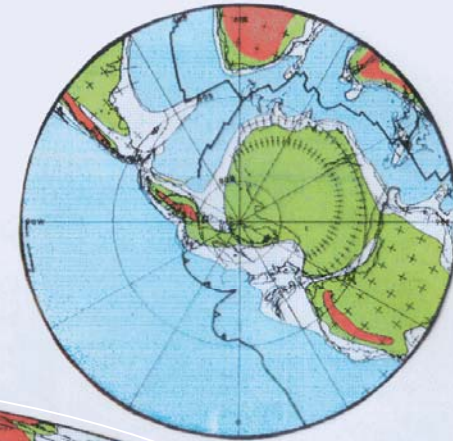


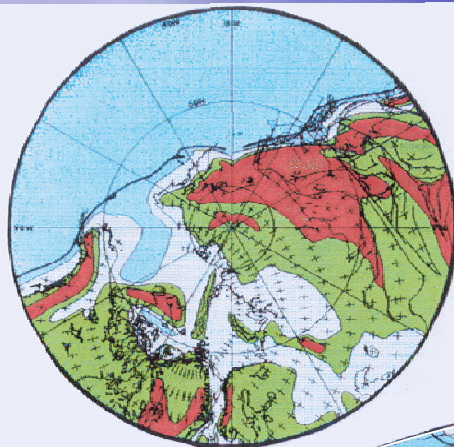
Middle Eocene
50.3 MA



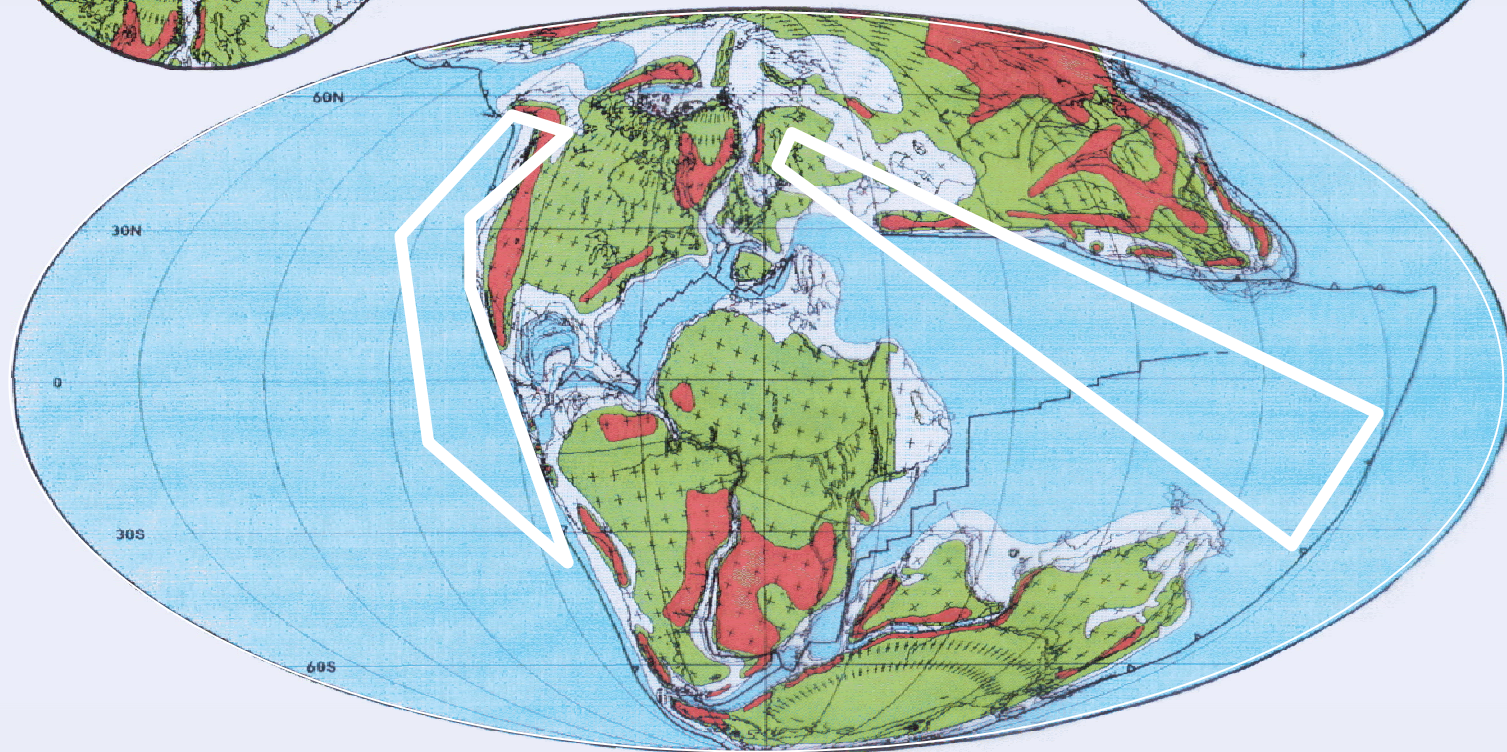
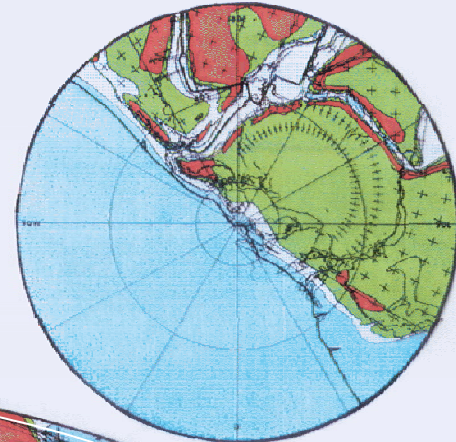


**Late Cretaceous
94.0 MA**





**Early Cretaceous
130.2 MA**



But what do we really see in the mantle?

True
velocity
anomaly

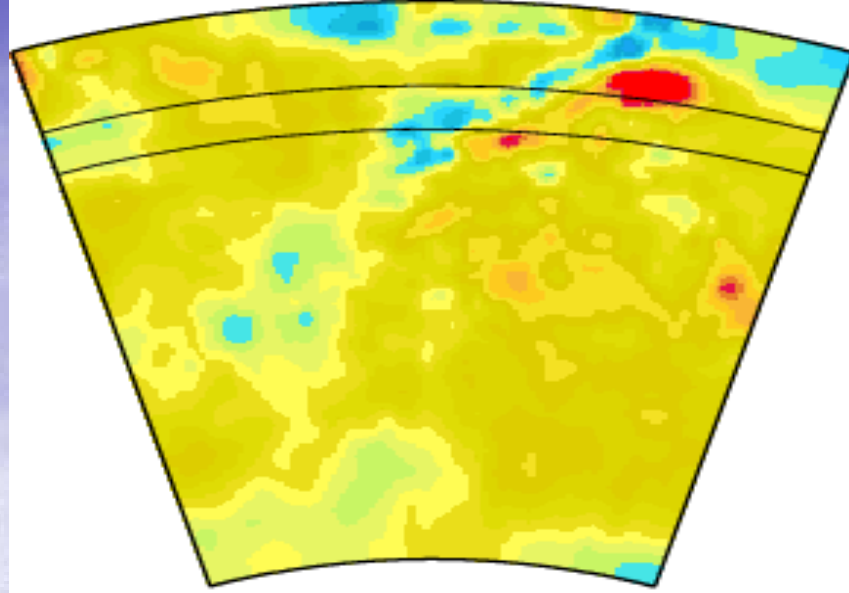
Ask Malcom
or Jeroen



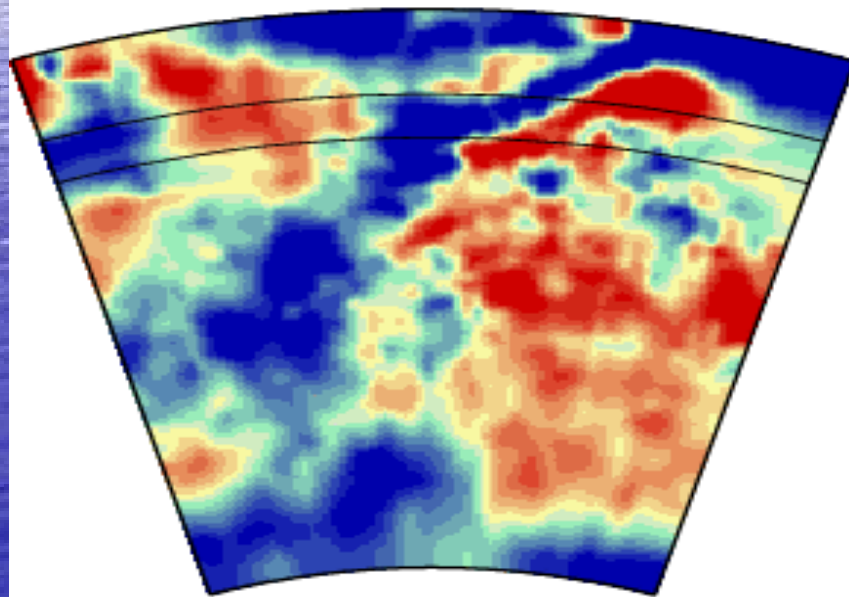
Inverted
velocity
anomaly

Physical
parameter

Ask a
mineralogist



Vmin Vmax



-0.5% 0.5%

Do we see
slabs in the
lower mantle?

Example of a priori **density** variations

A geodynamical model of temperature and pressure

A list of phase end-member thermo-elastic parameters

Olivine (α)	Forsterite	Mg_2SiO_4
$(\text{Mg,Fe})_2\text{SiO}_4$	Fayalite	Fe_2SiO_4
Olivine (β)	Mg-Wadsleyite	Mg_2SiO_4
$(\text{Mg,Fe})_2\text{SiO}_4$	Fe-Wadsleyite	Fe_2SiO_4
Ringwoodite (γ)	Mg-Ringwoodite	Mg_2SiO_4
$(\text{Mg,Fe})_2\text{SiO}_4$	Fe-Ringwoodite	Fe_2SiO_4
Magnesiowustite	Periclase	MgO
$(\text{Mg,Fe})\text{O}$	Wustite	FeO
Perovskite	Mg-Perovskite	MgSiO_3
(Mg,Fe,Al)	Fe-Perovskite	FeSiO_3
$(\text{Al,Si})\text{O}_3$	Al-Perovskite	Al_2O_3
Akimotoite	Mg-Akimotoite	MgSiO_3
$(\text{Mg,Fe})\text{SiO}_3$	Fe-Akimotoite	FeSiO_3
Orthopyroxene	Orthoenstatite	MgSiO_3
$(\text{Mg,Fe})\text{SiO}_3$	Orthoferrosilite	FeSiO_3
Clinopyroxene	Diopside	$\text{CaMgSi}_2\text{O}_6$
(Ca,Mg,Fe)	Hedenbergite	$\text{CaFeSi}_2\text{O}_6$
(Mg,Fe,Al)	Ca-Tschermak	$\text{CaAl}_2\text{SiO}_6$
$(\text{Al,Si})_2\text{O}_6$	Clinoenstatite	$\text{Mg}_2\text{Si}_2\text{O}_6$
	Clinoferrosilite	$\text{Fe}_2\text{Si}_2\text{O}_6$
Grenat	Pyrope	$\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$
$(\text{Ca,Mg,Fe})_3$	Almandin	$\text{Fe}_3\text{Al}_2\text{Si}_3\text{O}_{12}$
$(\text{Mg,Fe,Al,Si})_2$	Grossulaire	$\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$
Si_2O_7	Mg-Majorite	$\text{Mg}_4\text{Si}_4\text{O}_{12}$
	Fe-Majorite	$\text{Fe}_4\text{Si}_4\text{O}_{12}$

Composition

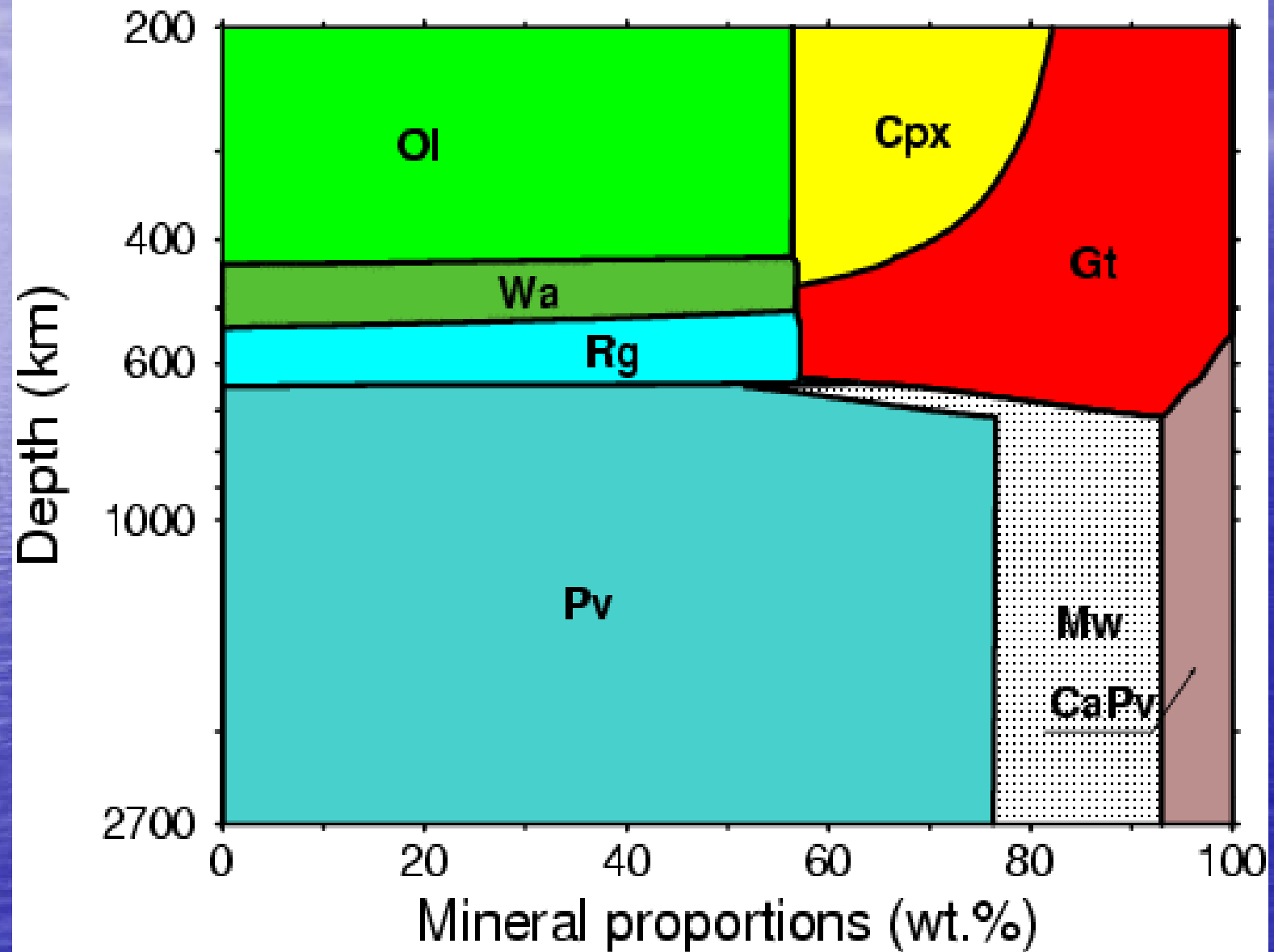
Si, O

ization

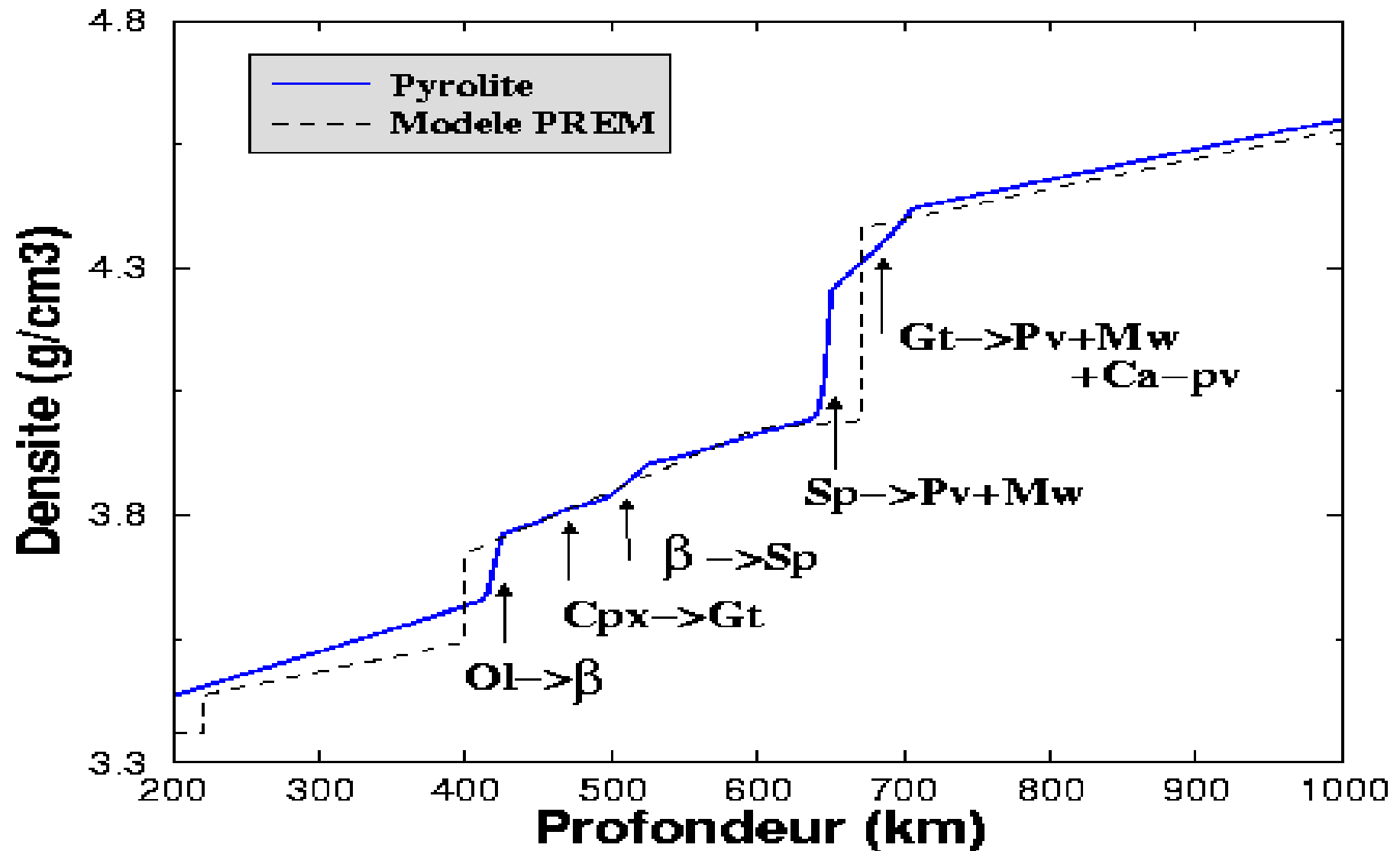
pattern...)

neralogy

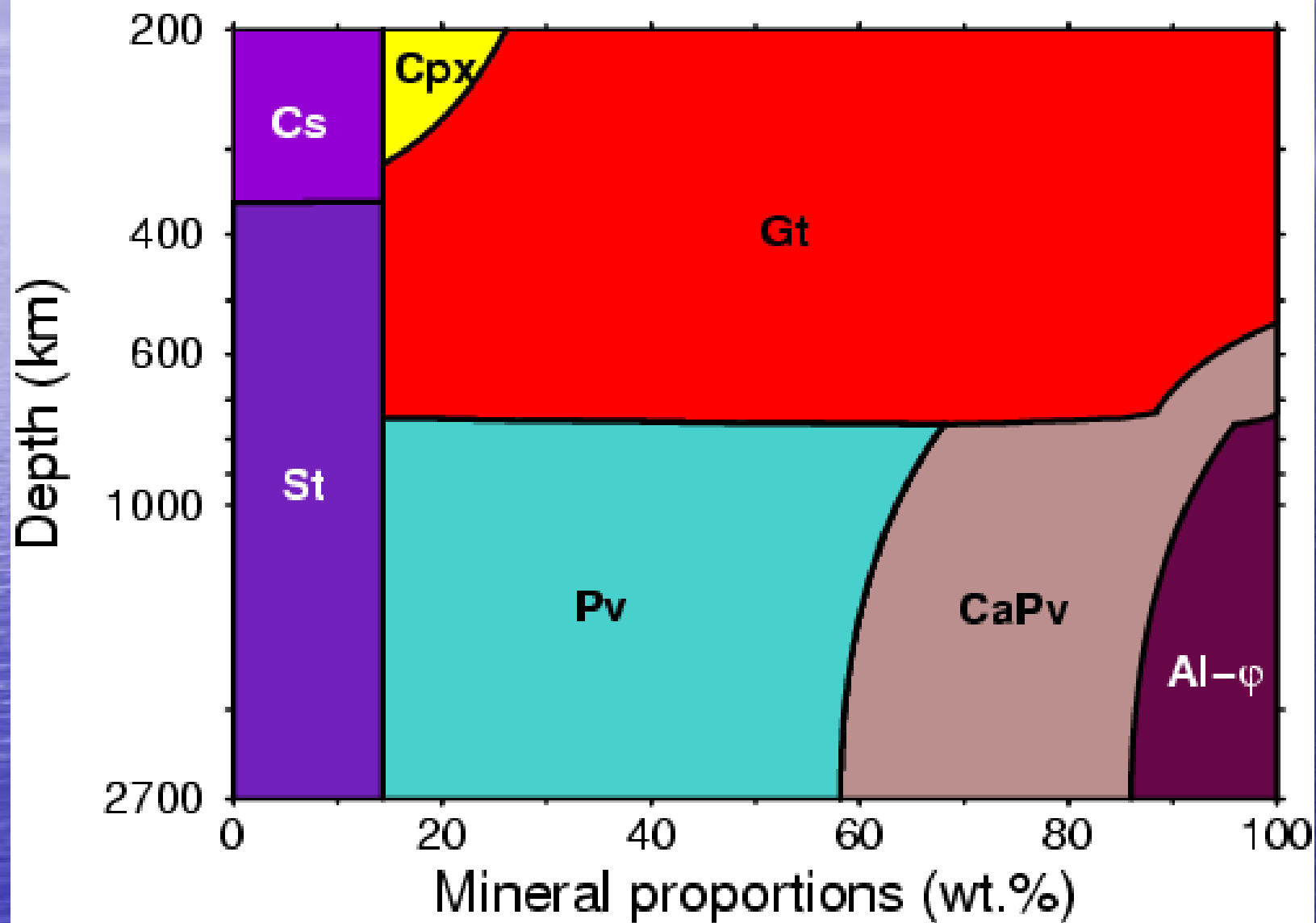
Pyrolite

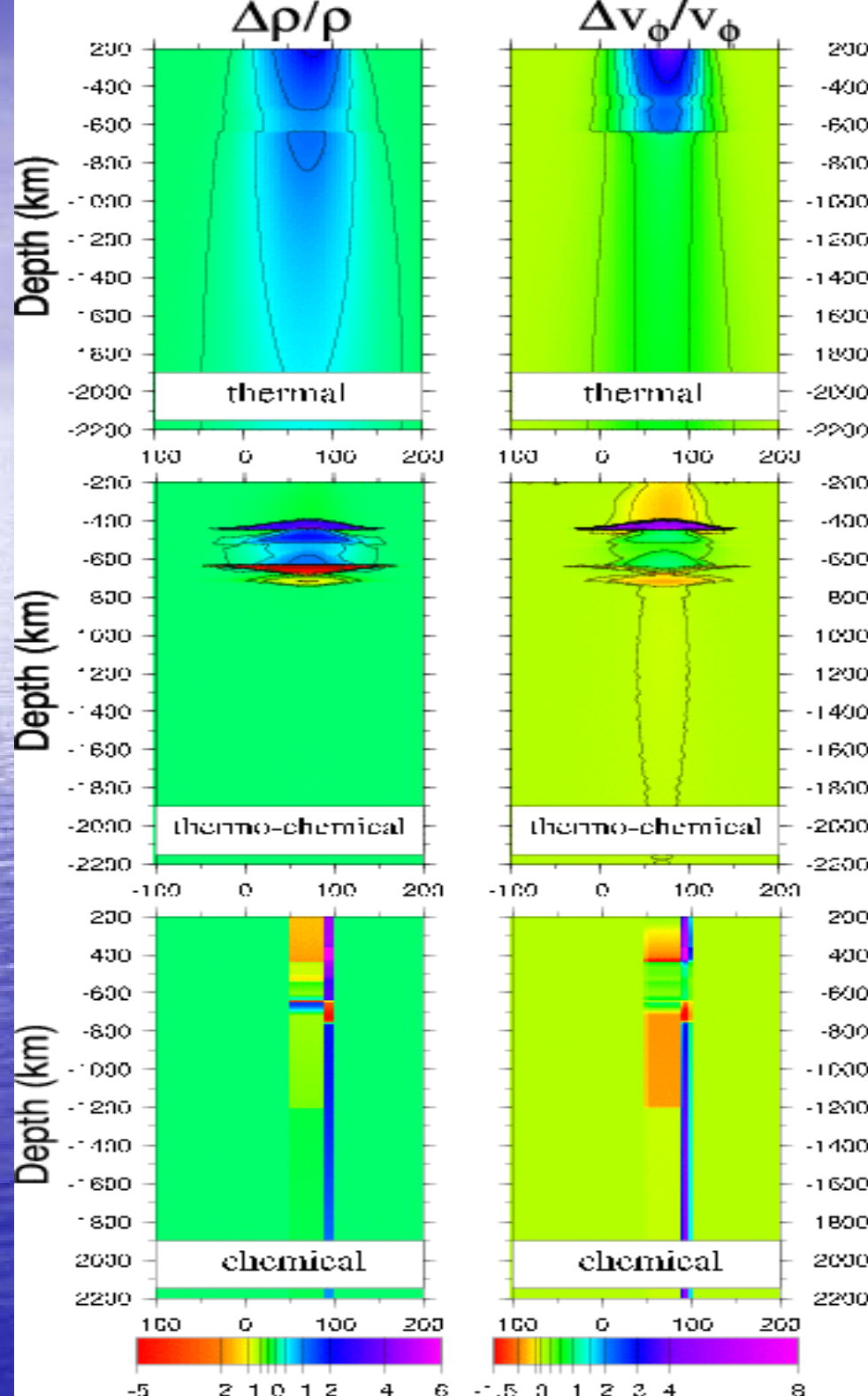


Mineralogy vs. Seismology



MORB





Example of a vertically
sinking slab:

Basalt+harzburgite+pyrolite

Intrinsic thermal effect

+

Changes in mineralogy
to maintain equilibrium

+

Changes in mineralogy
according to changes in
elemental composition

- The thermal anomalies,
 - the thermo-chemical anomalies,
 - the compositional anomalies
- have similar amplitudes...

The end of purely thermal convection models?...

Ra=1000.000000 Ras=3000.000000 Numx = 257 Numy = 129

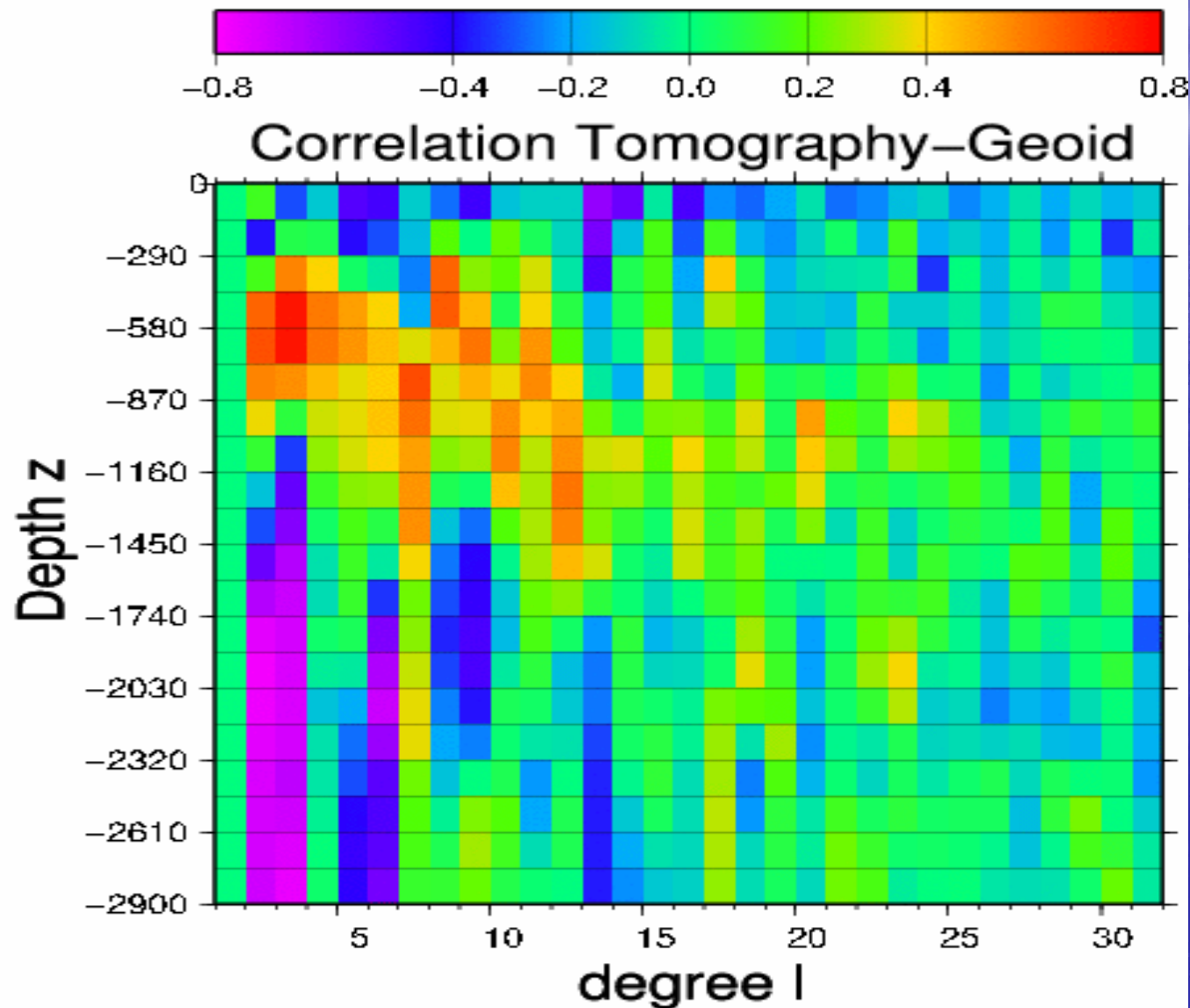
Temp Step: 1
Time: 0.000000
Min: 0.
Max: 1.

Chem
Min: 0.000000
Max: 2.667106

Uran min/max 0.00 / 5.00
Min: 1.849999
Max: 1.850000

He4 min/max 0.00 / 5.00
Min: 3.699998
Max: 3.700001

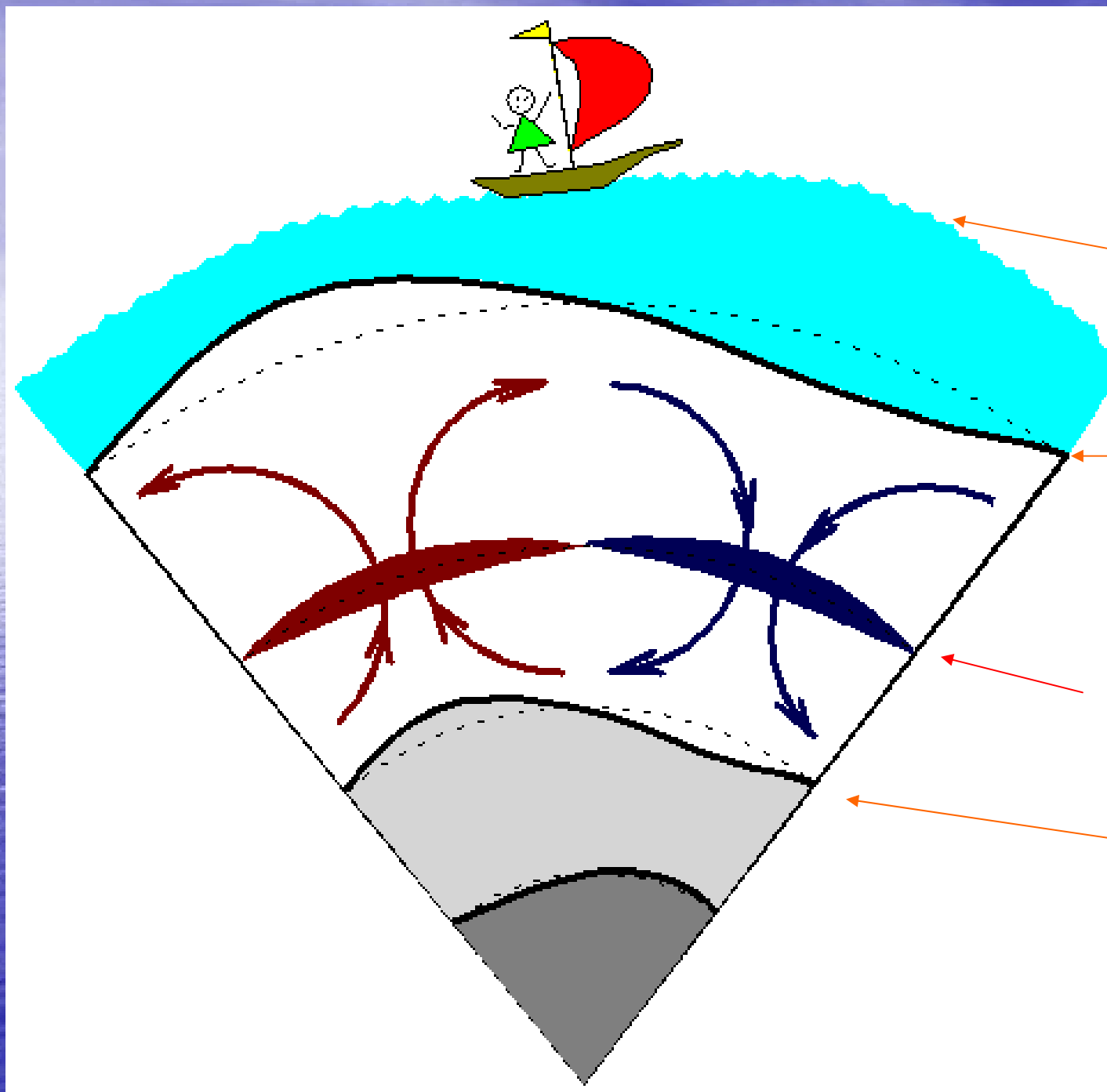
2. Rheology



Negatif

Positif

Negatif



Geoid

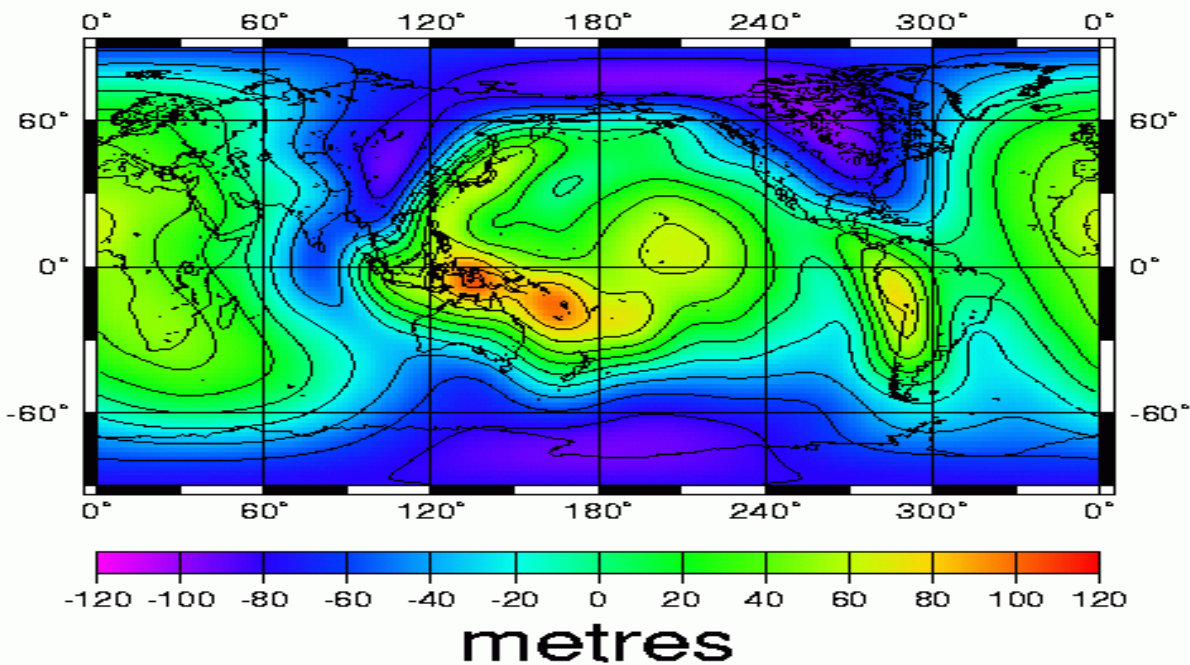
Deflexion
“Dynamic
Topography”

Loads

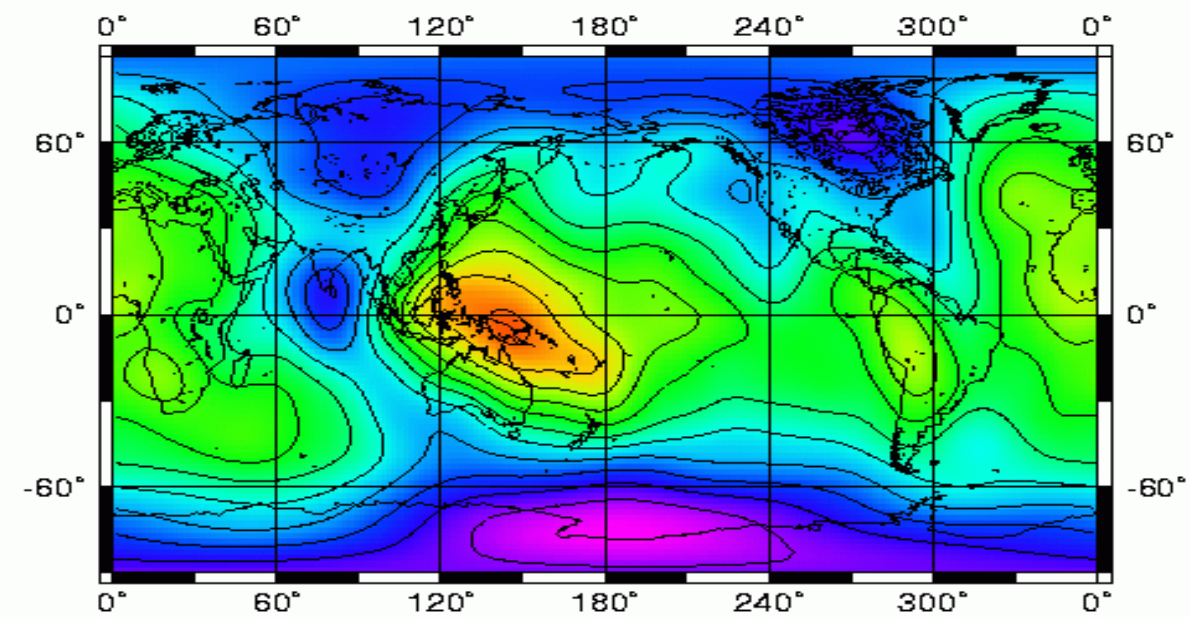
Deflexion

The gravity is anti-correlated with the internal density, except in the vicinity of a factor 10-100 viscosity increase...

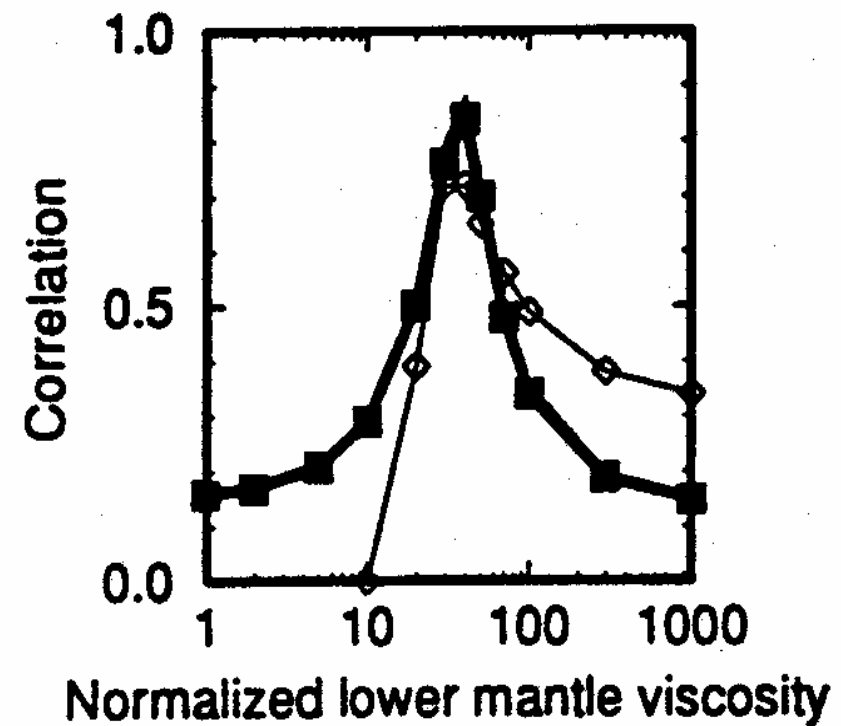
GEOIDE SYNTH.



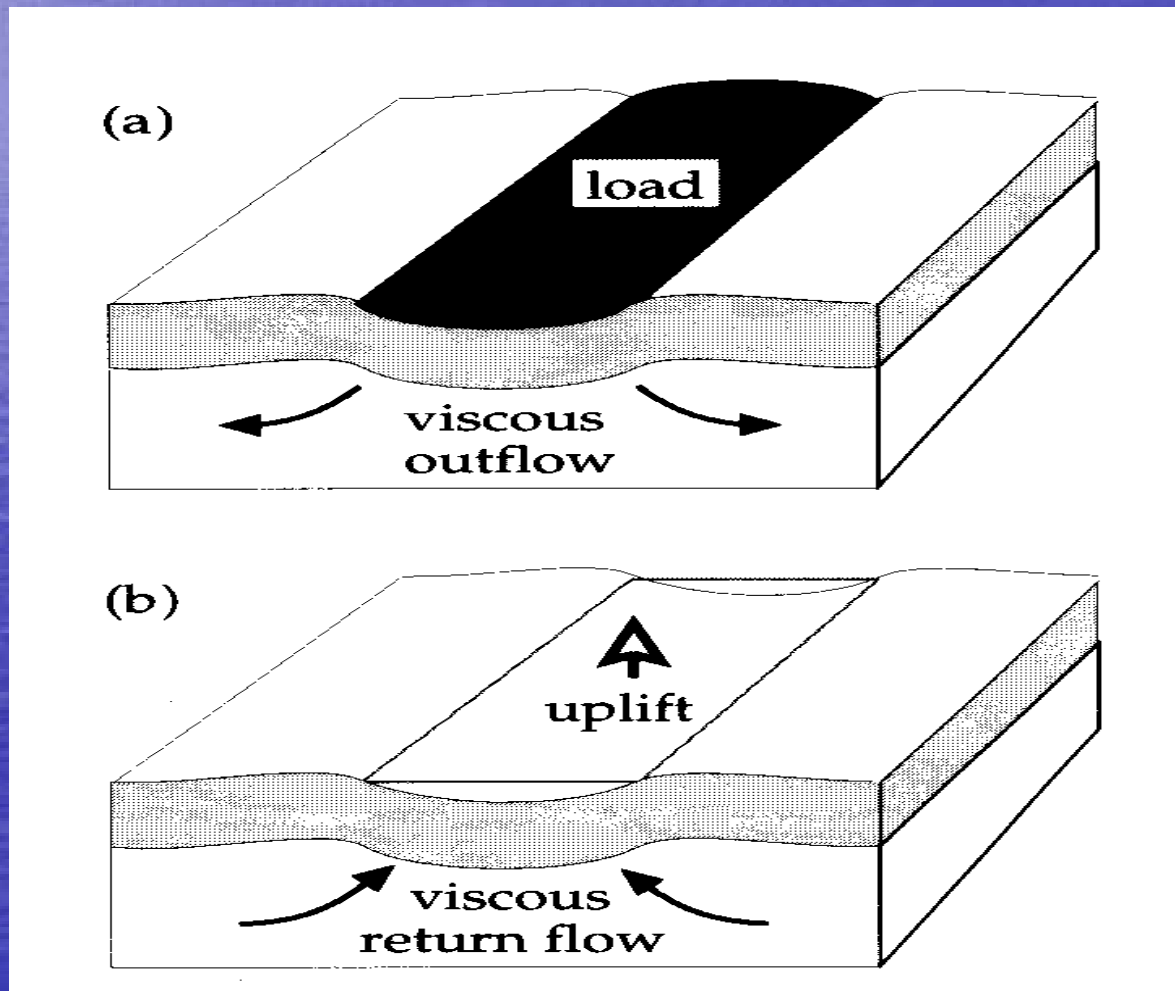
GEOIDE



COMP. VS. OBS. GEOID



...which is in agreement(?) with post-glacial rebound modeling...



Kurt...

Finland



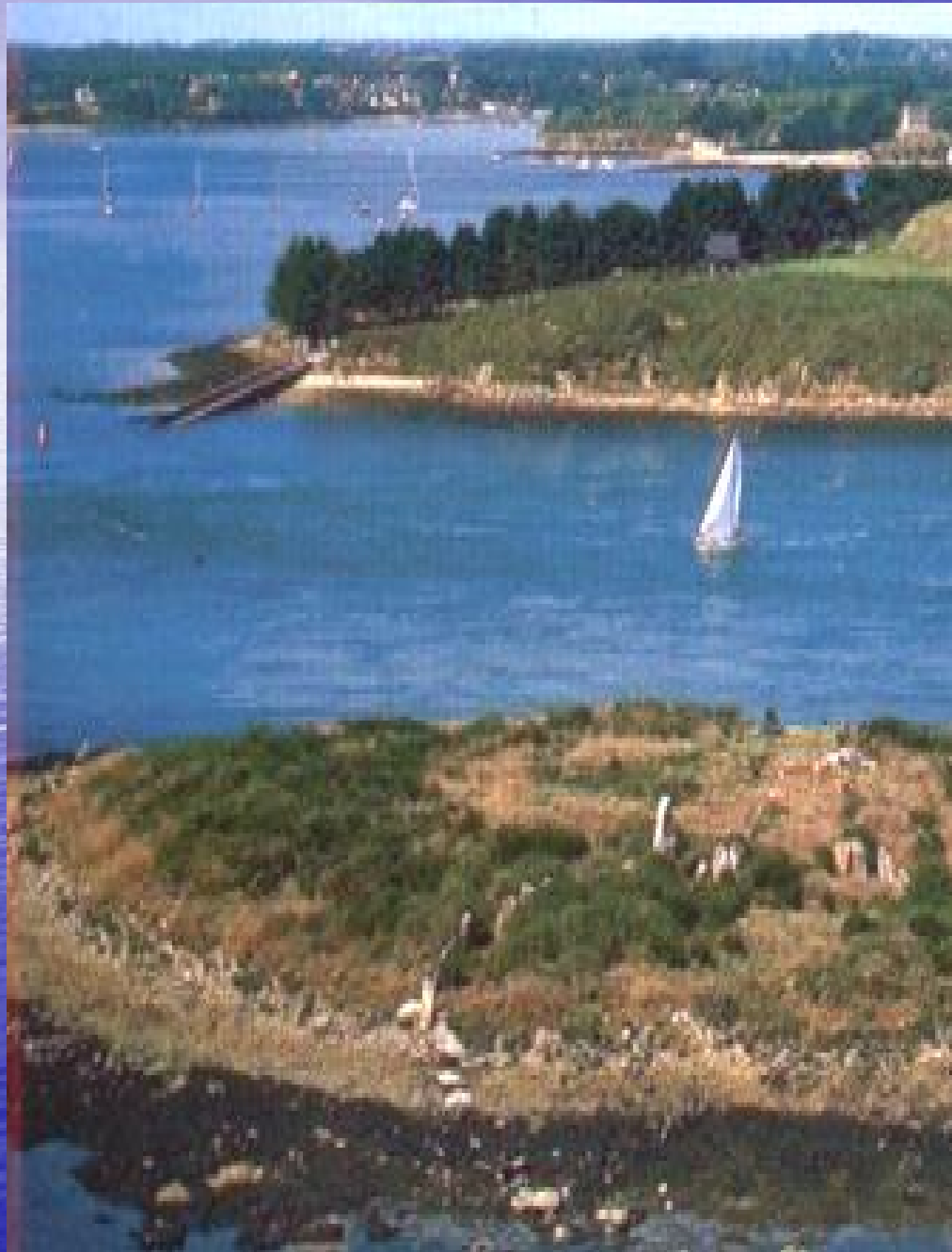
Hudson



Scotland



Bretagne

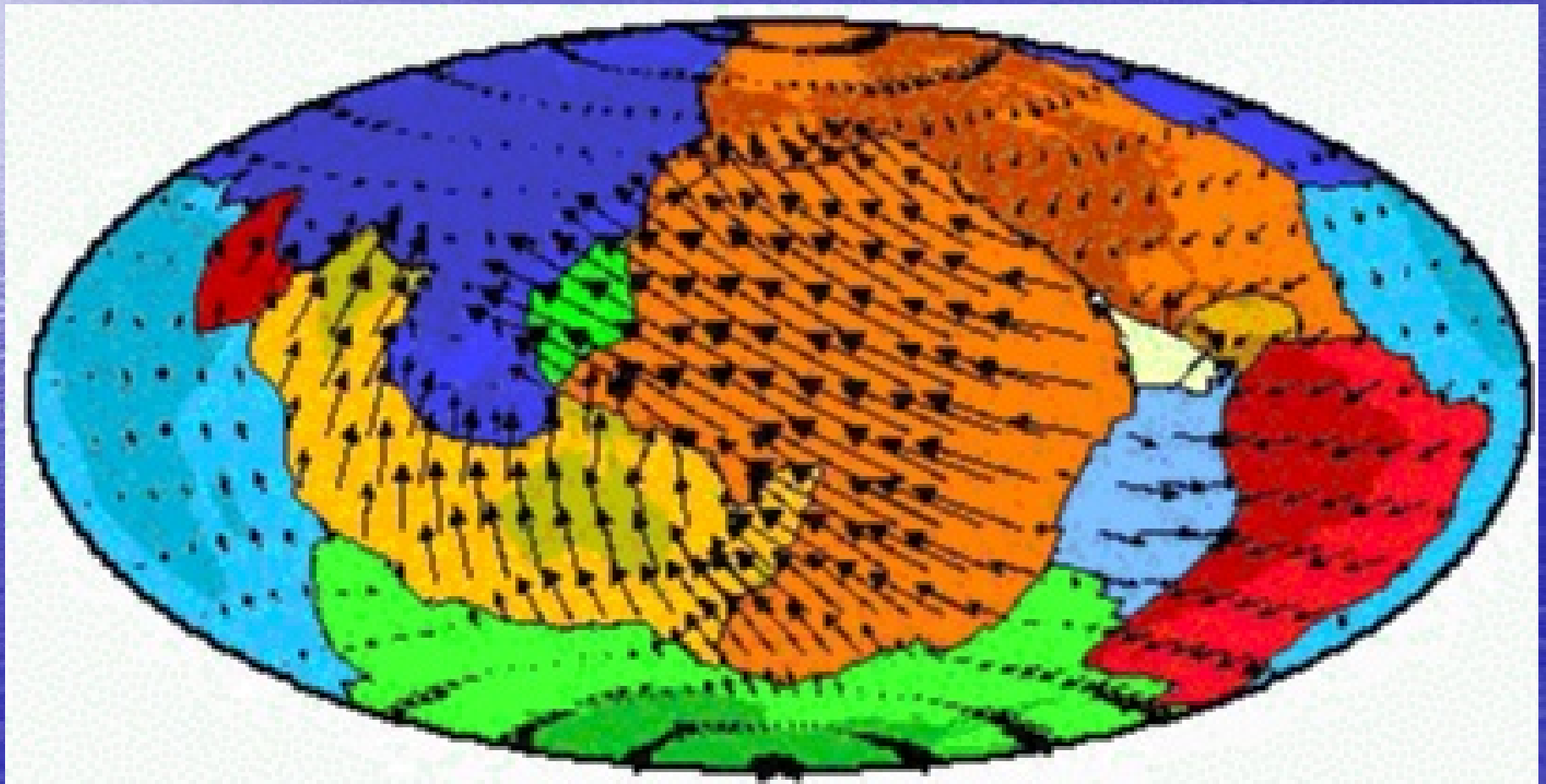


- But do we see the dynamic topography?
not a lot!
- But should the rheology be only linear visco-elastic, depth-dependent?
no 1)

Rheology depends on stress, temperature, melt, water content, grain size, composition...

Patrick, Greg, Mark...

- But should the rheology be only depth dependent?
no 2)

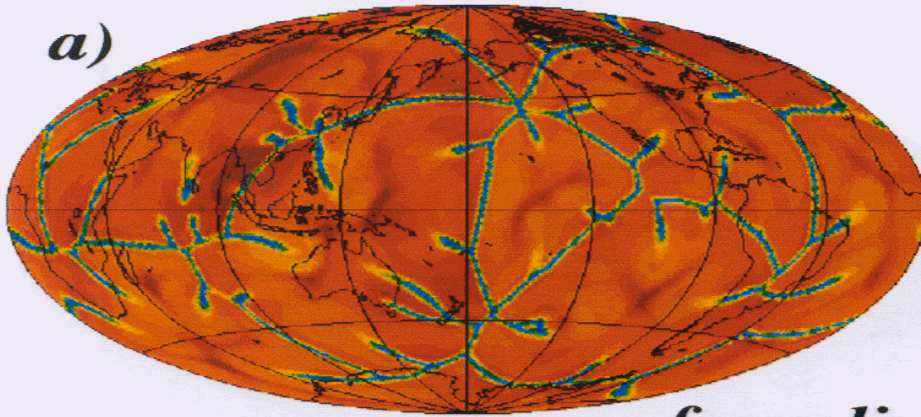


layered viscosity convection

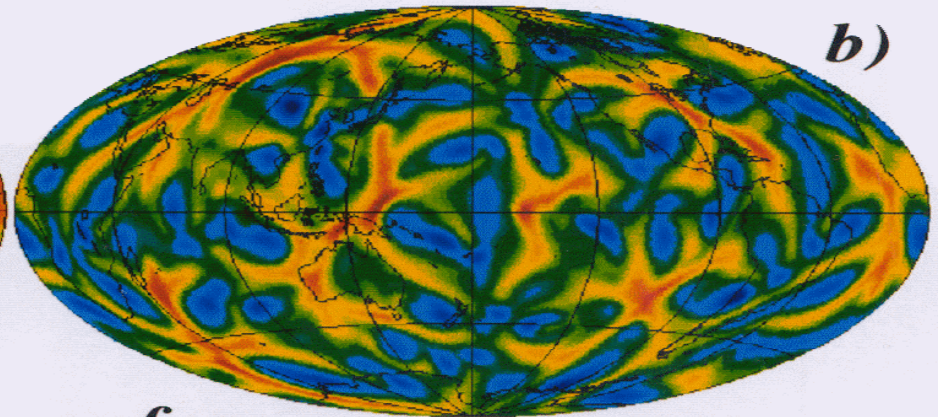
upper mantle

mid mantle

a)

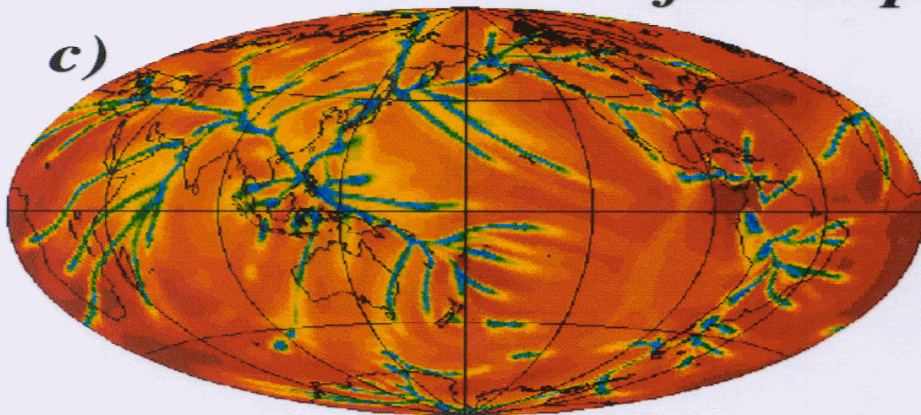


b)

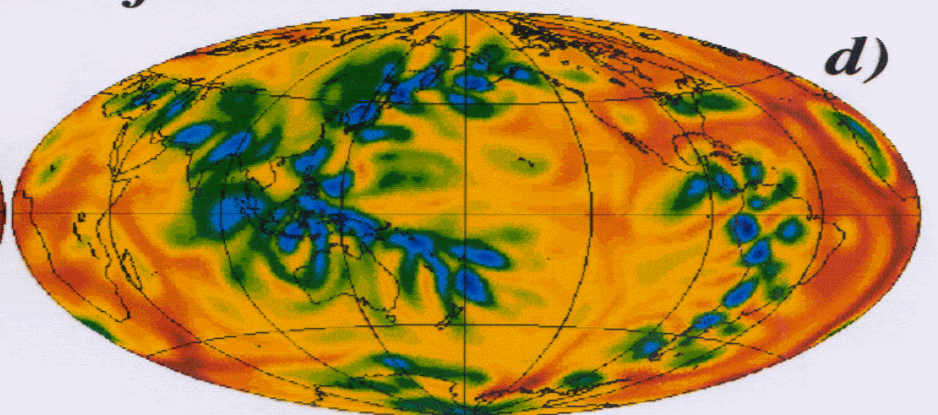


free slip surface

c)

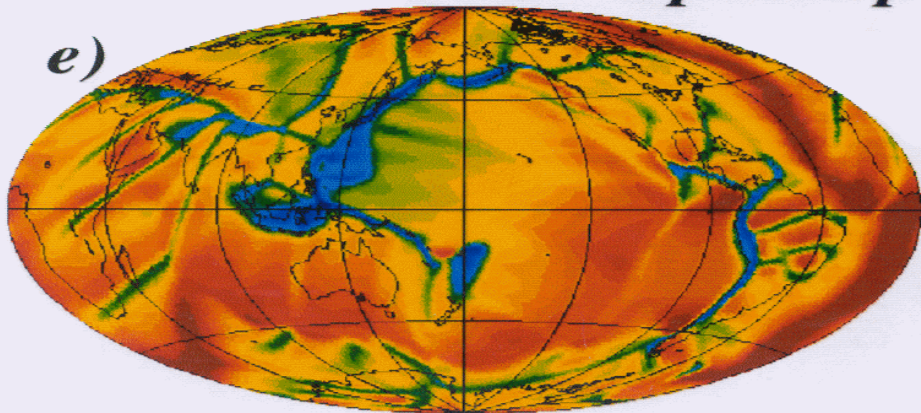


d)

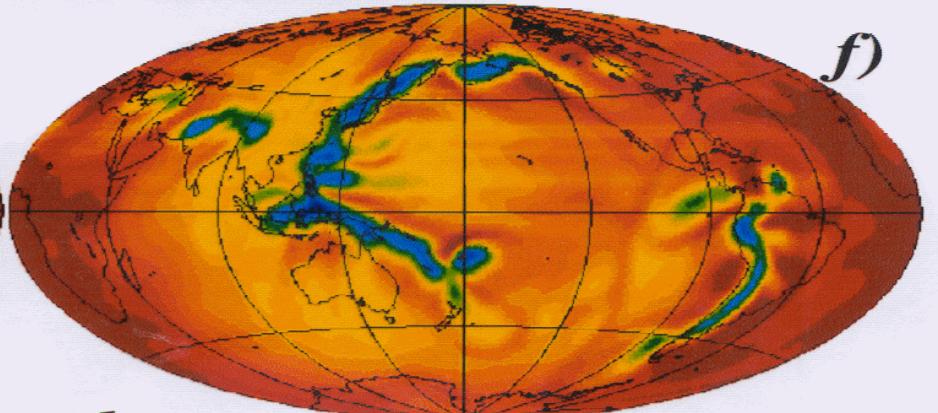


+ imposed plate motions

e)

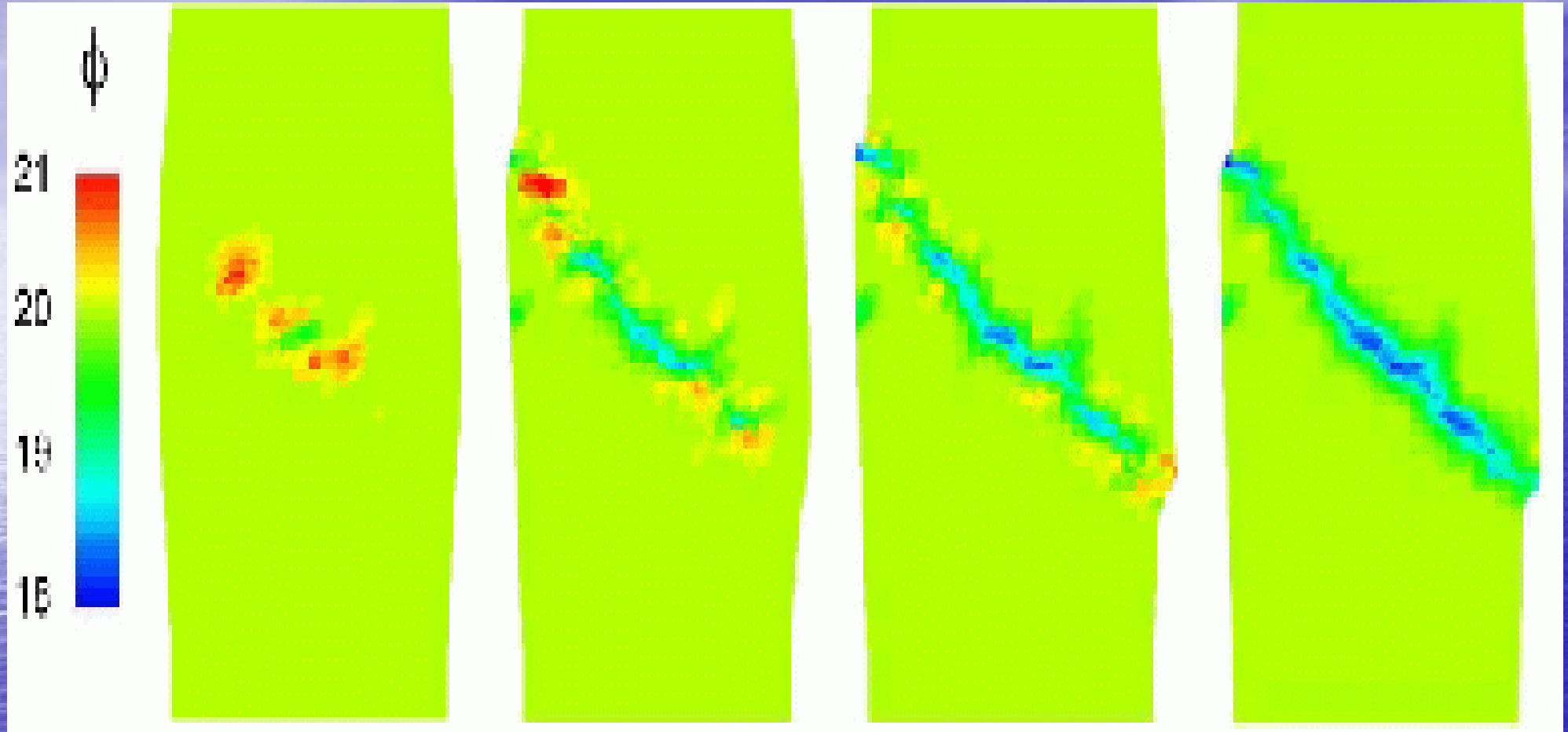


f)



+ strong plates

LOCALISATION

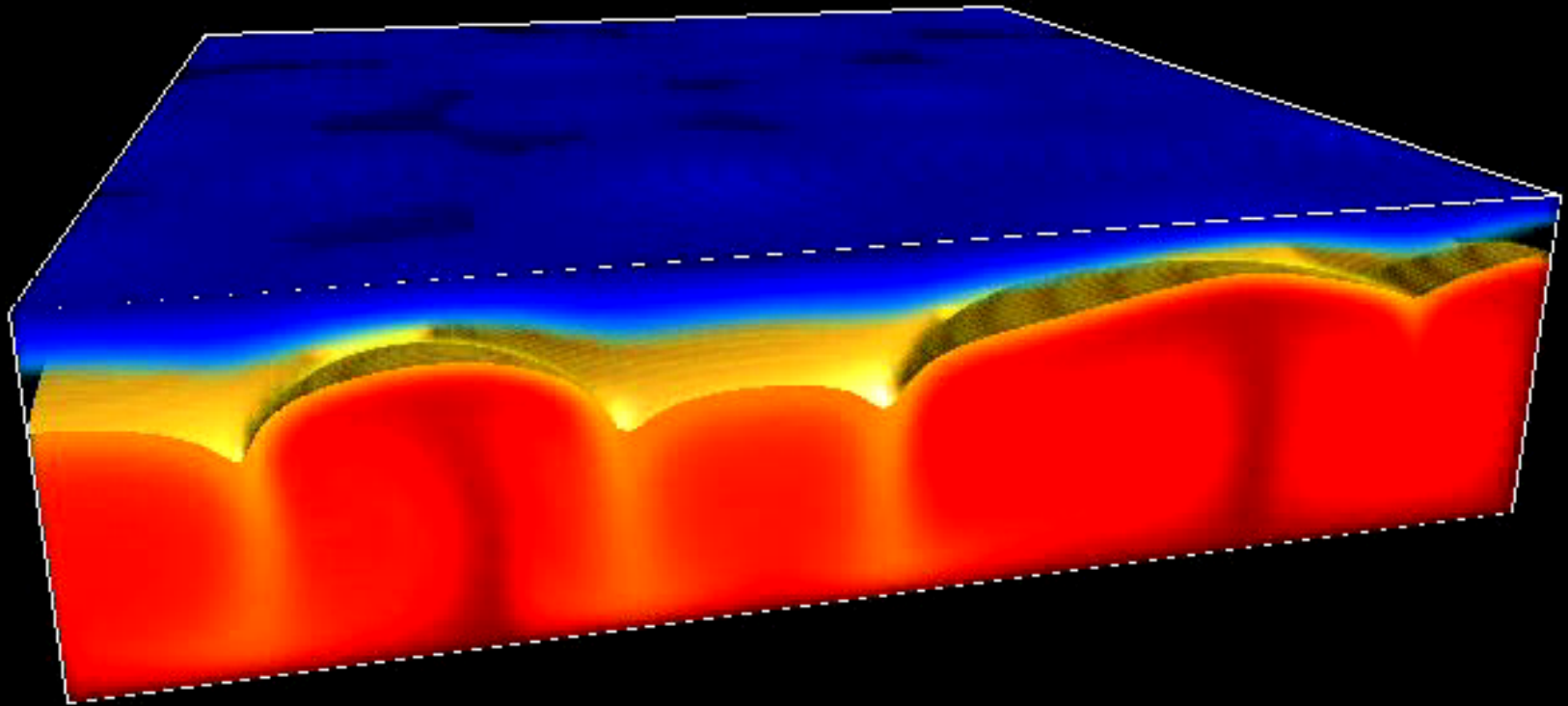


The rheology has a memory.

It depends on ~~stress~~, ~~temperature~~, melt, water content, grain size, composition...

Nstep = 1220

Time = 0.118323



Joerg

Damage variable x

$$\text{Rheology} = F(x)$$

$$Dx/dt = \text{damage-healing}$$

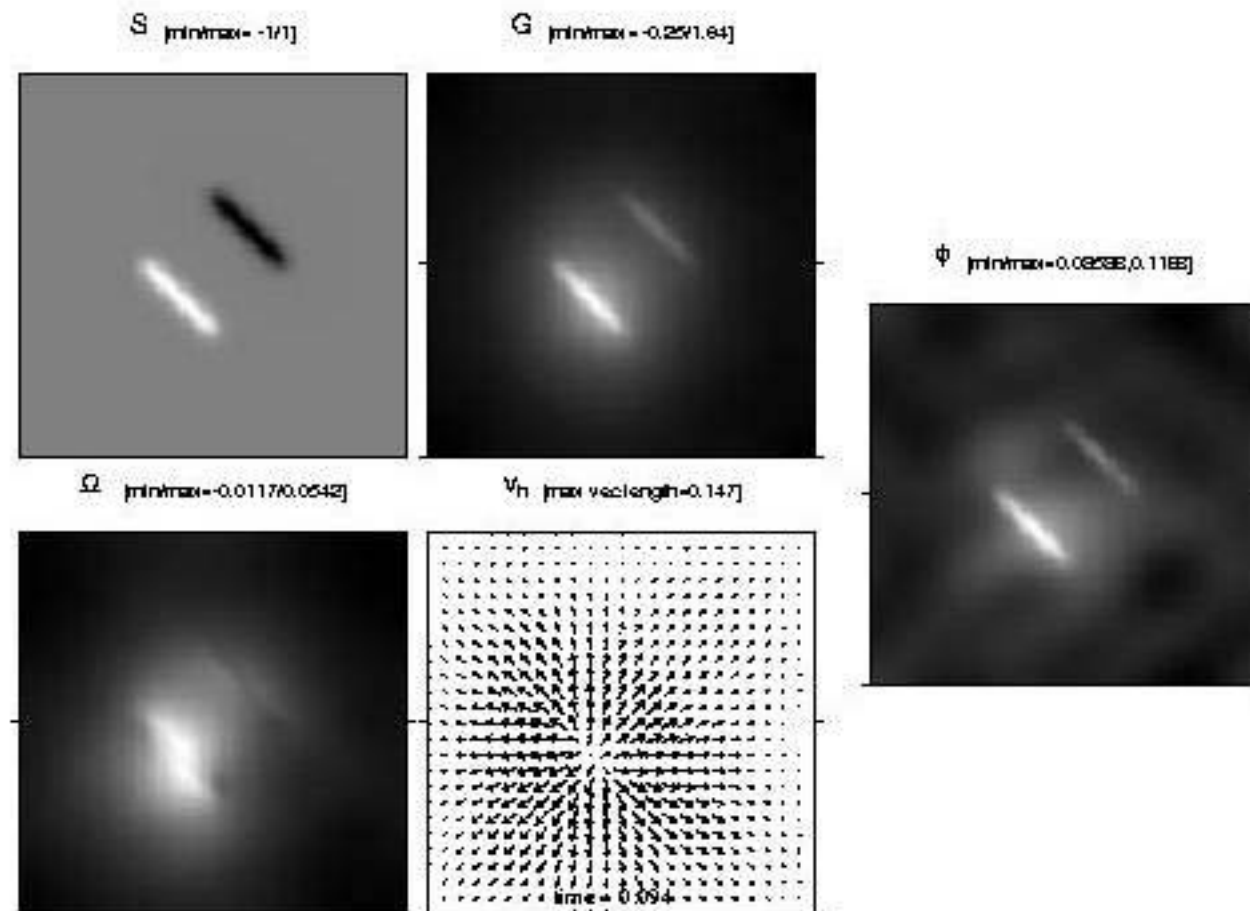
Ex:

$$DT/Dt = \text{shear_heating-diffusion}$$

$$D \langle \text{grain size} \rangle / Dt = \text{fracture_nucleation-coarsening}$$

$$D \text{ porosity} / Dt = \text{microcracking-compaction}$$

Porosity Related Damage $f_{\alpha_0} = 0, f_{\phi} \neq 0$



Grain Size Related Damage $f_{\alpha_0} \neq 0, f_{\phi} = 0$

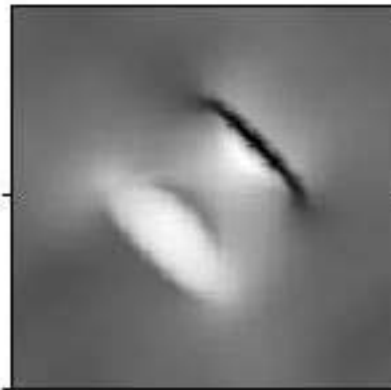
S [min/max = -1/1]



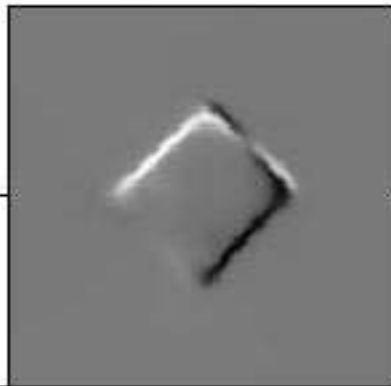
G [min/max = -0.0108/0.00723]



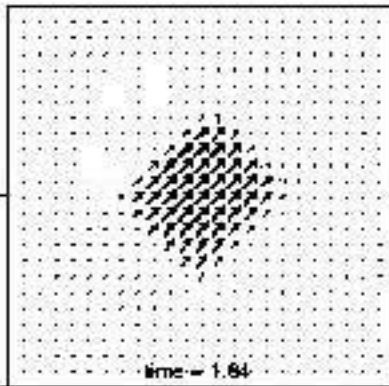
Φ [min/max = 0.0891/0.0893]



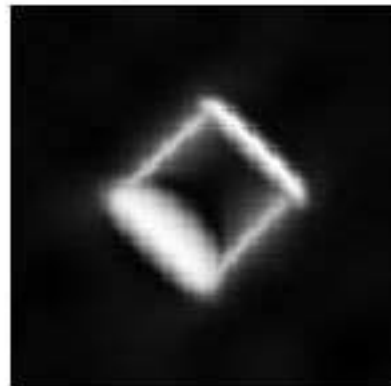
Ω [min/max = -1.58/1.74]



V_h [max vec length = 0.108]



\mathcal{R} [min/max = 0.9824/1.339]



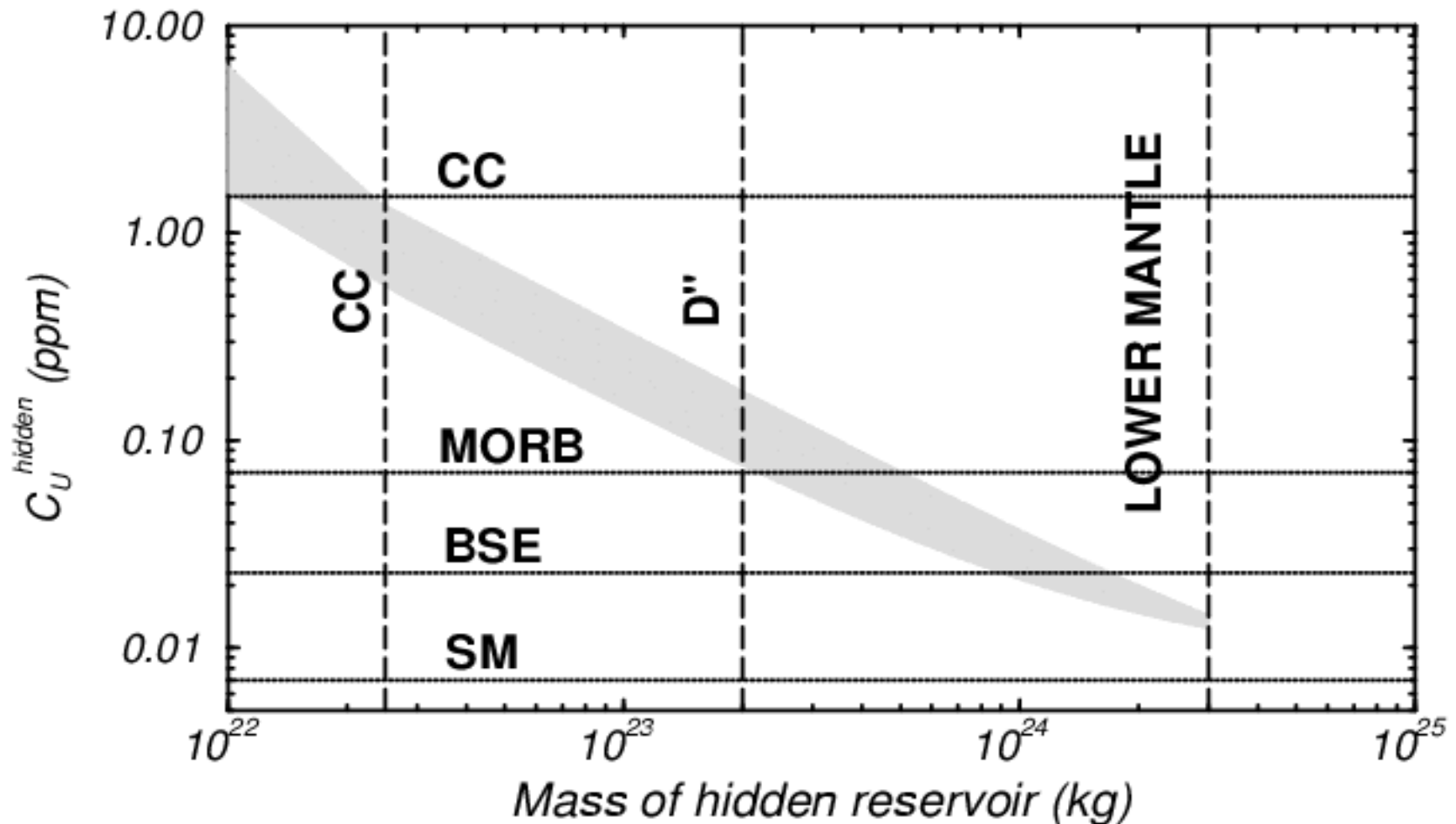
The end of models with linear depth
dependent models...



Mixing/geochemistry

Mass Balance for trace elements

Bulk Sil. Earth = Crust+Morb source+Hidden res



There is a hidden reservoir...

- ✓ Slightly depleted=lower mantle
- ✓ Primitive=50% of the mantle
- ✓ With higher concentration of incompatible elements than BSE
(D'' with oceanic crust composition)

^{40}Ar gon

Produced in the Earth 940 pmol/g

Atmosphere 44%

Crust 3.5%

Upper mantle .9 % (25 pmol/g)

Lower mantle 52 % (720 pmol/g)

But K/U??

50-200 pmol/g

Another K-rich
reservoir?

Total Heat Sources U, Th, K...

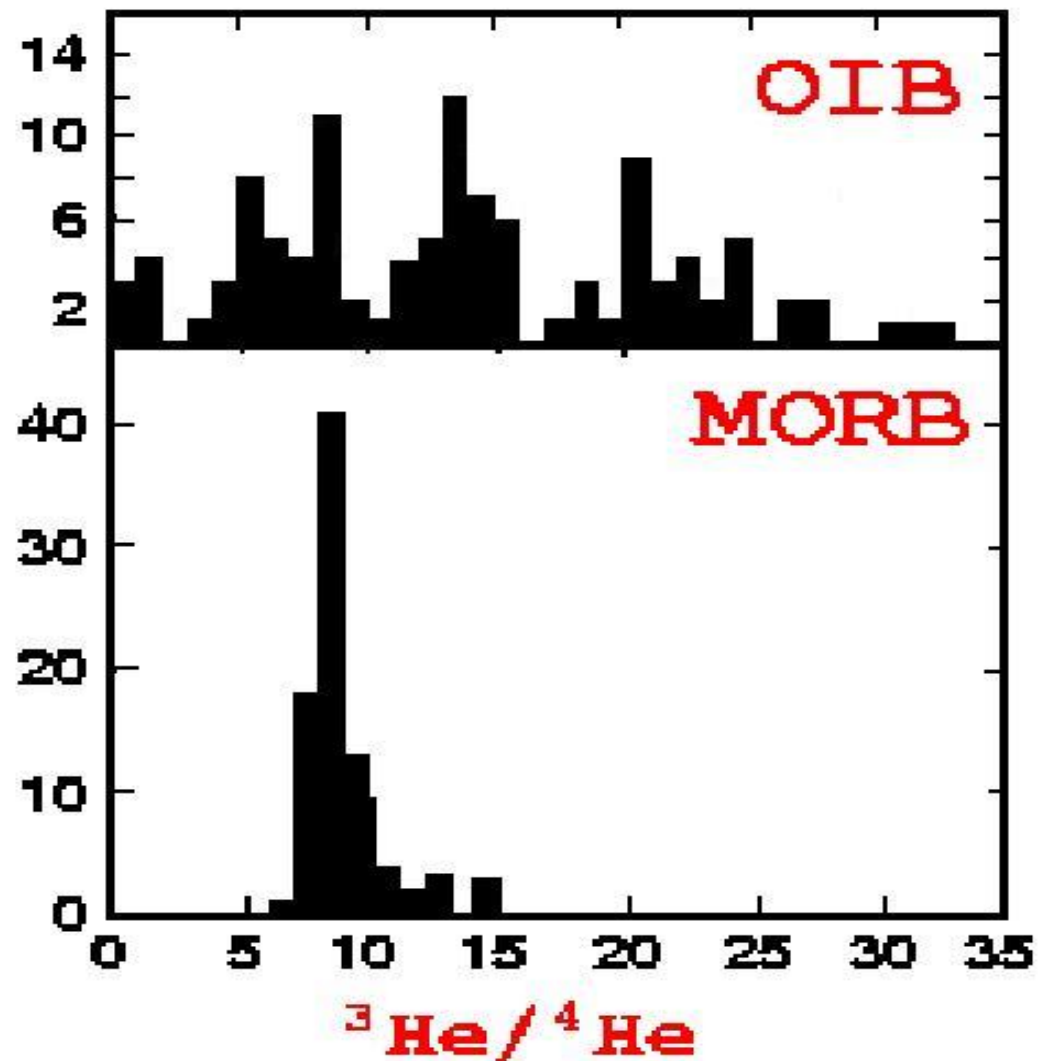
$H = 20 \text{ TW}$

$Q = 43 \text{ TW}$

Cooling = Production - Surface heat loss

Cooling 23 TW !

Rare Gas



✓OIBs are more heterogeneous than MORBs (but there is a continuous range between OIBs and MORBs)

Erasing heterogeneities

Mixing = stirring + diffusion

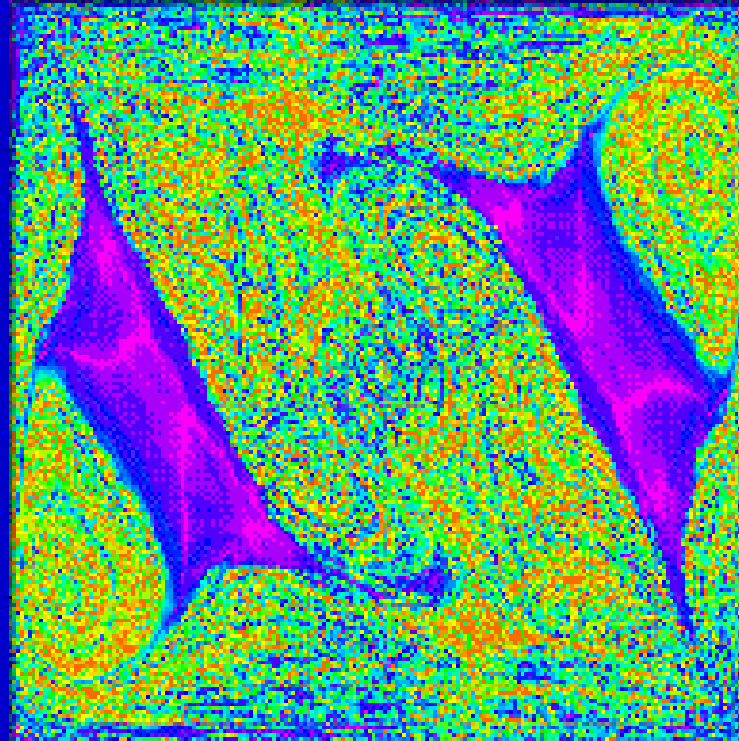
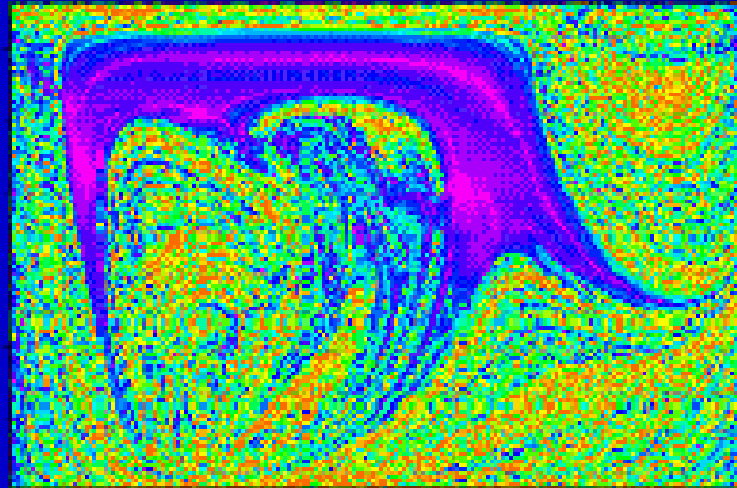
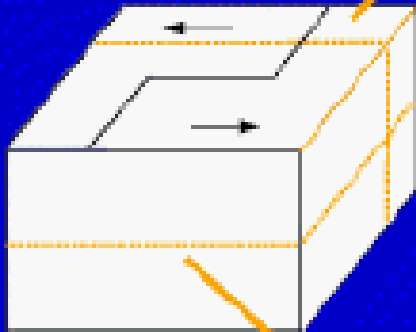
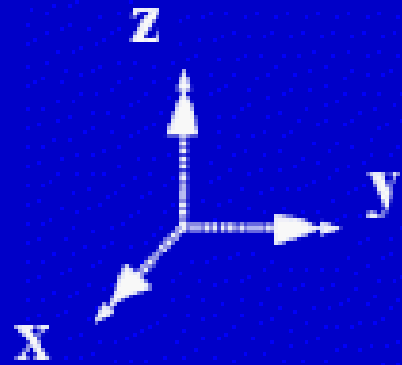
Diffusion... $L^2 = 2 D t$

$$\text{Log}(\epsilon' L^2/D)/2\epsilon'$$

advected

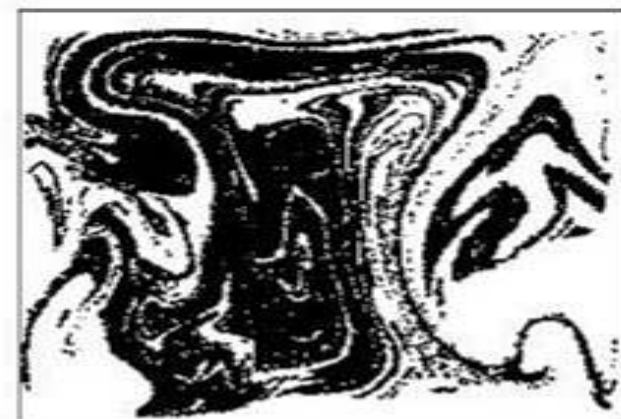
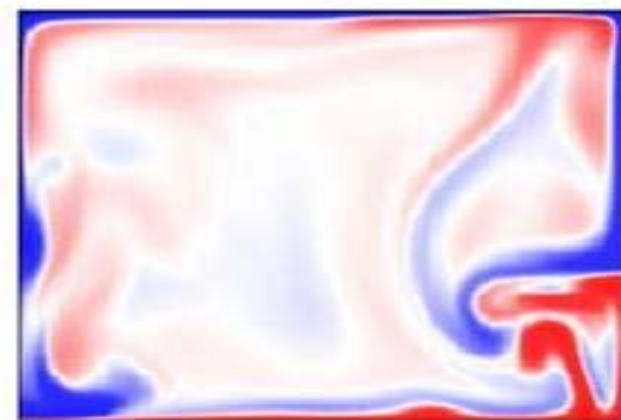
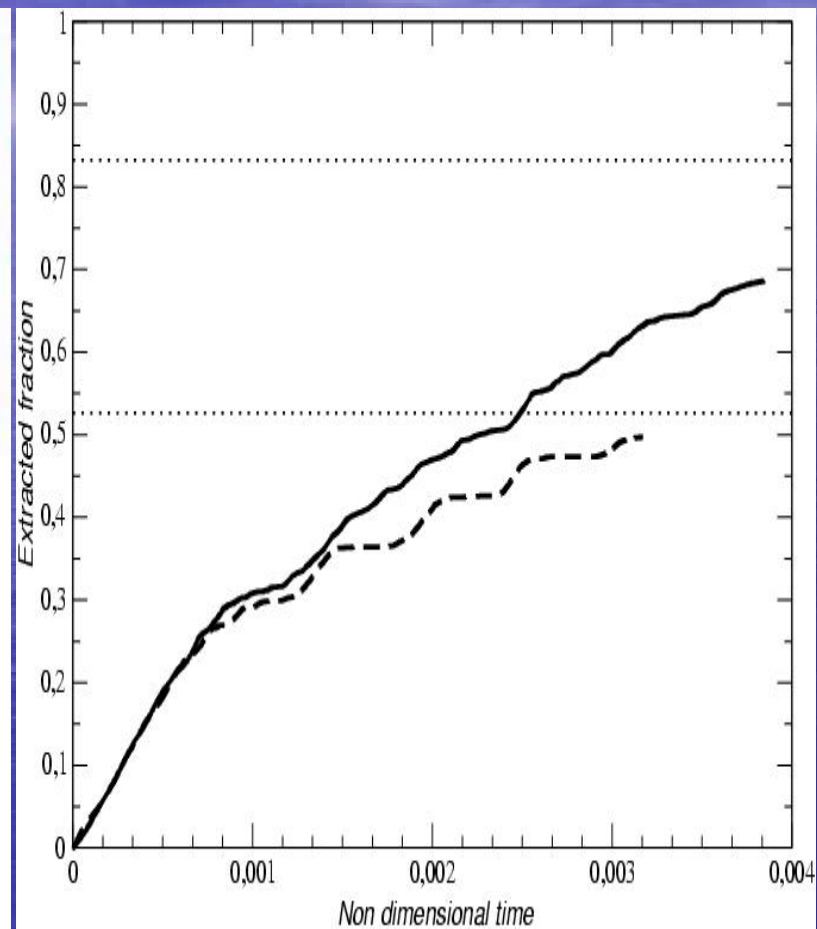
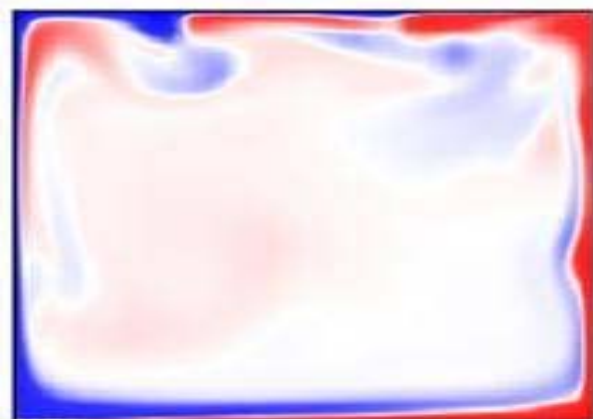
totally mixed

- ✓ Erasing U in the oceanic crust:
2 byrs (?)
- ✓ Erasing He in the oceanic crust:
1 byr (?)



Steady
convection

An efficient mixing is not necessarily
due to a strong temporal variability



In whole mantle convection models
when the volume of remaining
primitive material decreases, its
probability to be processed under a
ridge also decreases

3-20 % of primitive mantle
(filaments, lumps...)

Various MORB and OIB recipes

Stratified Cake:

MORB are from the SM (shallow mantle)

OIB are a mixture between SM and primitive mantle

Marble Cake:

The whole mantle is made of recycled oceanic crust + peridotite (recycled lithosphere+primitive mantle)

The difference between MORB and OIB are mostly due to mixing differences

Various MORB and OIB recipes

Cake with raisins:

MORB are from the dough

OIB taste from the raisins

Water filter:

Fast (OIB) and slow upwellings are filtered by water extraction at 400 km.

PRIMITIVE MANTLE

ORIGIN OF ANOMALIES

Stratified cake	LARGE QUANTITY	INITIAL
Marble Cake	SMALL QUANTITY	CONSEQUENCE OF TECTONICS
Cake with raisins	SMALL QUANTITY	INITIAL
Water filter	SMALL QUANTITY	CONSEQUENCE OF TECTONICS

SLABS IN THE LM

DIFF. LM/UM

Stratified cake	NO	YES
Marble Cake	YES	NO
Cake with raisins	YES	NO
Water filter	YES	YES

Still a lot to do....

The possibility of confronting seismic images, mineralogic and geochemical data with geodynamic models is still rather new

Knowing the limits of our knowledge is a first step to have new ideas...